



HELLENIC REPUBLIC
National and Kapodistrian
University of Athens

Formation of large-scale structure by turbulence in planetary atmospheres

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Physics Department

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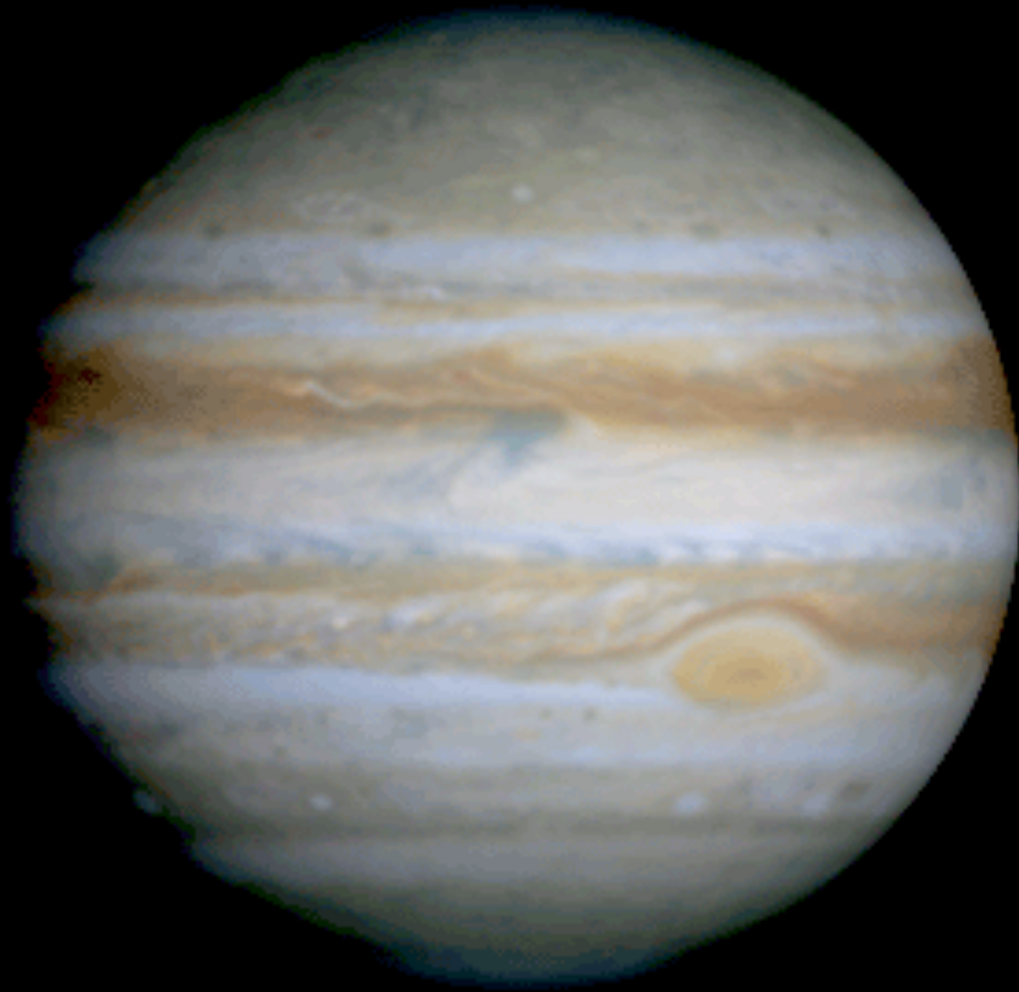
structure of talk

- ▶ introduction to the physical problem
- ▶ formulation of the theory (S3T)
- ▶ study of the stability of homogeneous turbulent state
- ▶ comparison of S3T predictions with direct numerical simulations and verification of the theory
- ▶ stability of inhomogeneous turbulent states & relation with jet mergers
- ▶ relation of modulational instability of Rossby waves with S3T instability of homogeneous state (if time allows)
- ▶ summary

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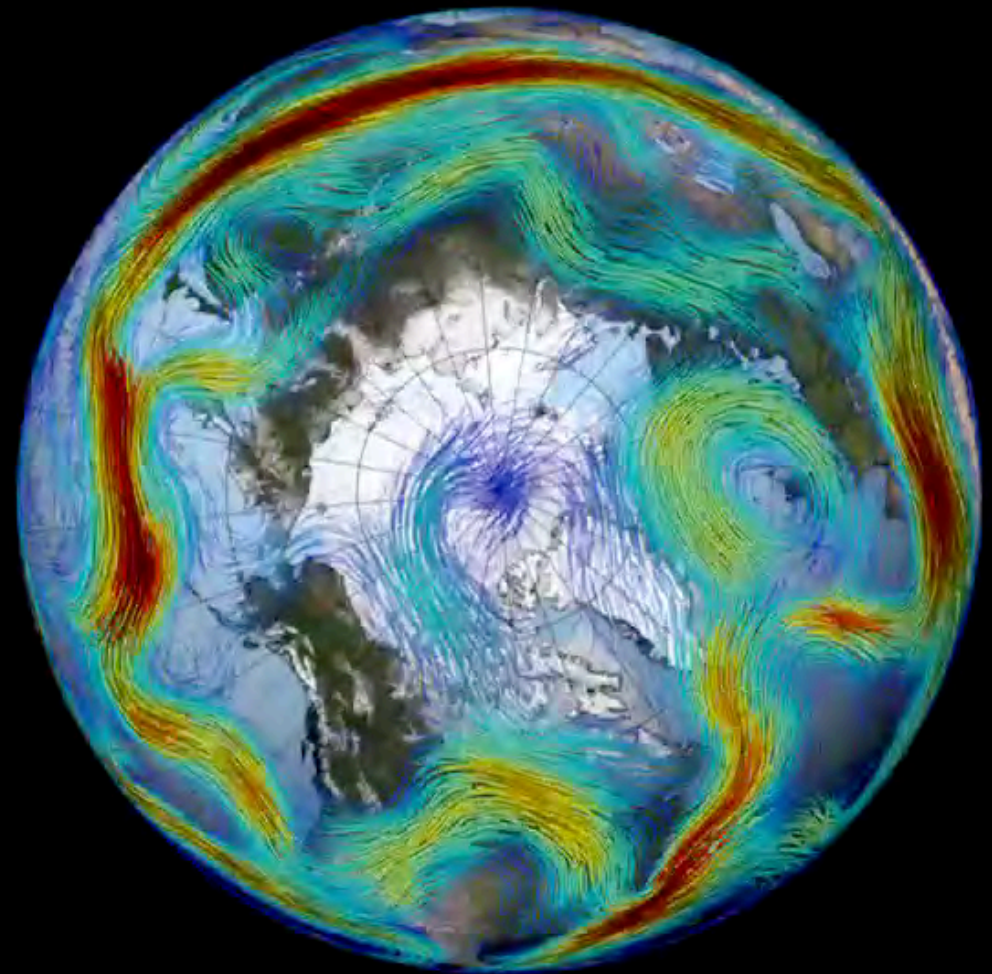
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Planetary turbulence is anisotropic and inhomogeneous I



banded Jovian jets

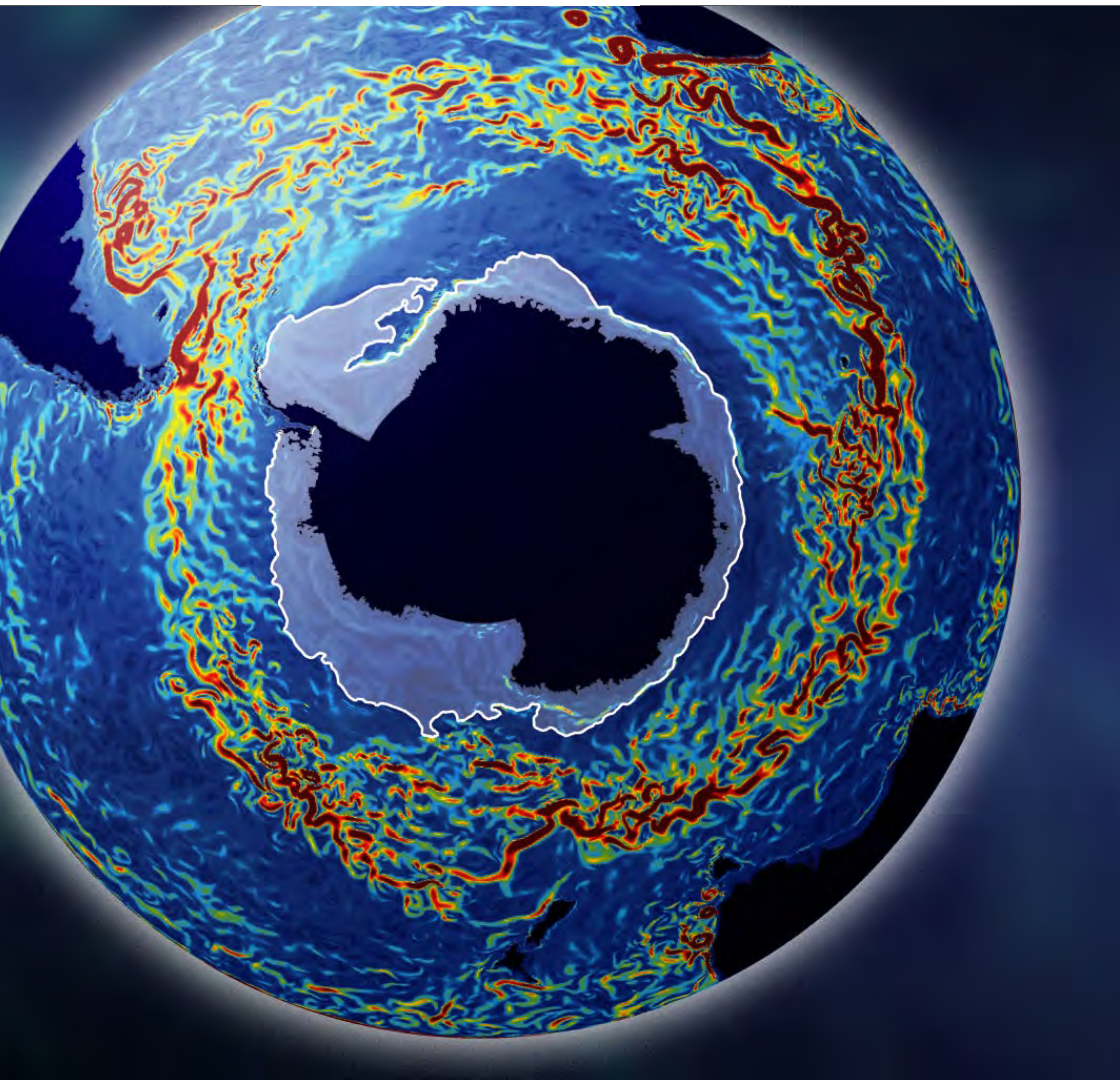
NASA/Cassini Jupiter Images



polar front jet

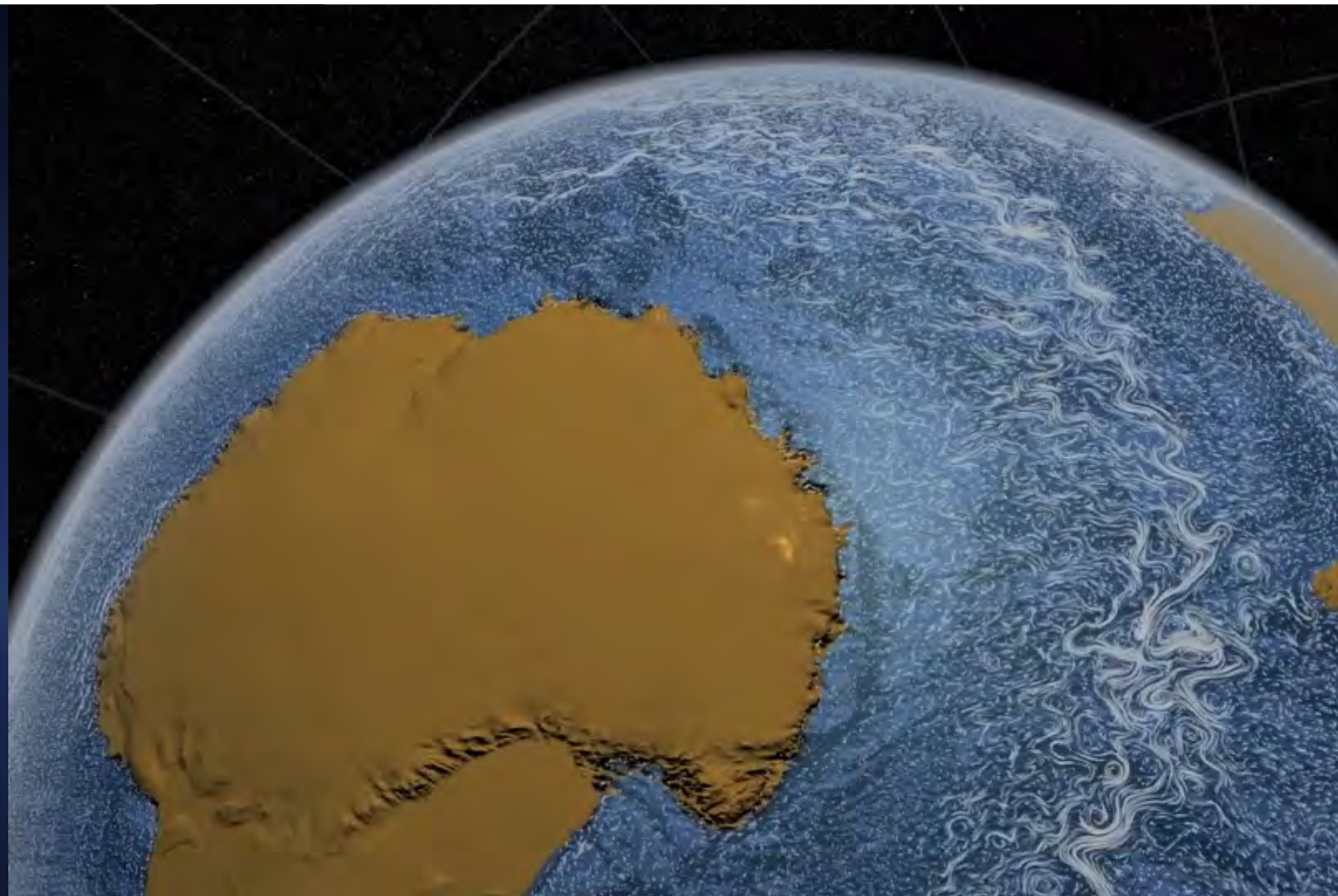
NASA/Goddard Space Flight Center

Planetary turbulence is anisotropic and inhomogeneous II



computer simulation

San Diego Supercomputer Center, UCSD



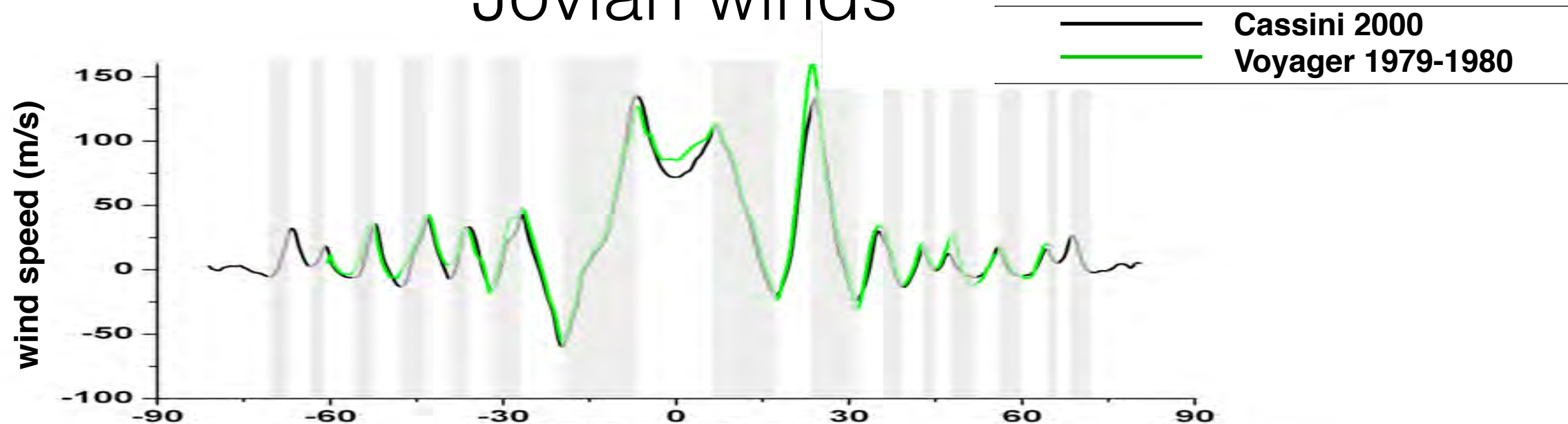
ACC

satellite observations

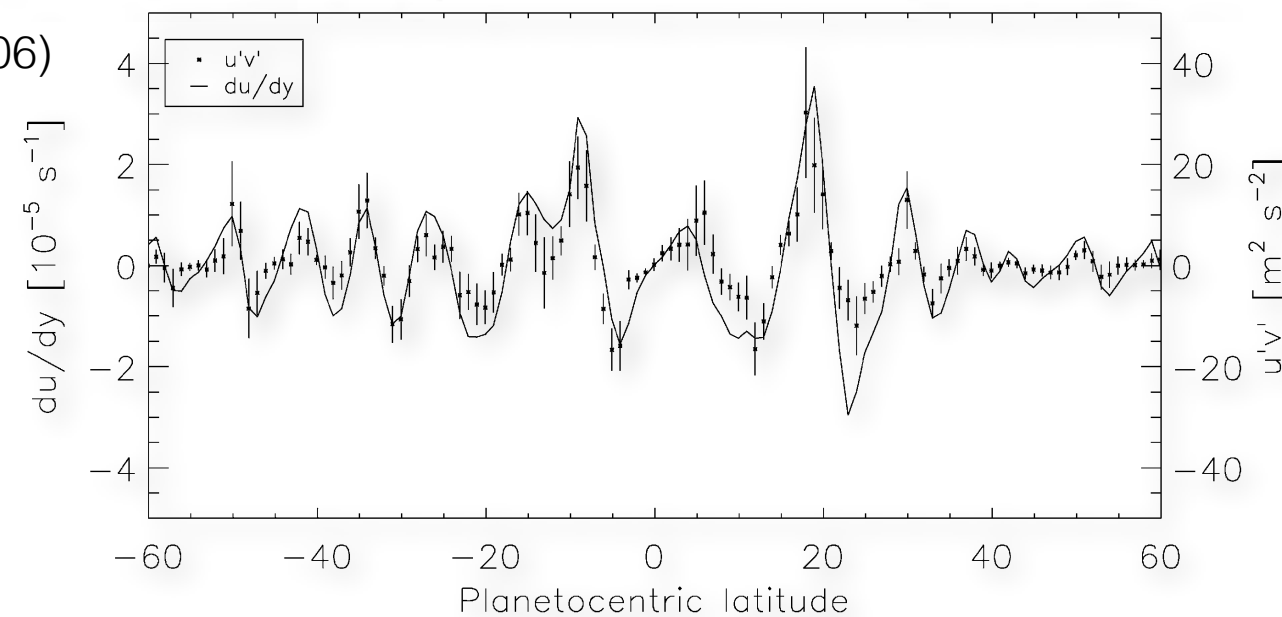
NASA/Goddard Space Flight Center

Jets appear “steady” and are eddy-driven

Jovian winds



(Salyk et. al. 2006)

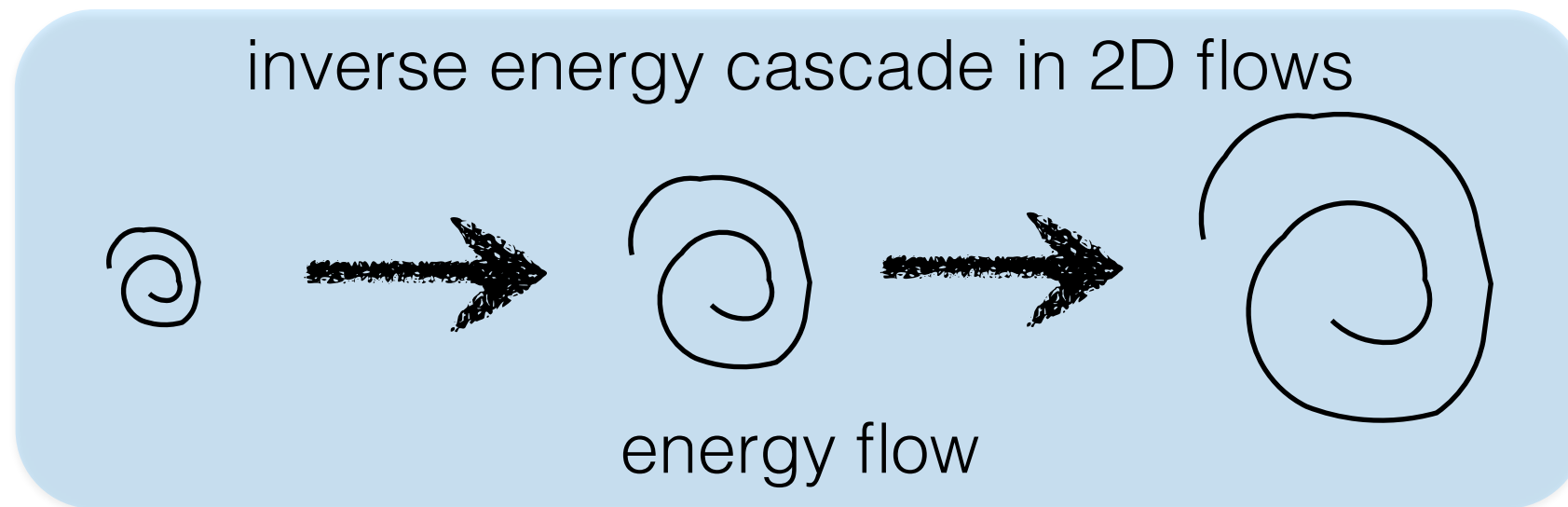


$$\overline{u'v'} = \kappa \frac{\partial U}{\partial y}$$

$$\kappa \approx 10^6 \text{ m}^2 \text{ s}^{-1}$$

$$\partial_t U = -\partial_y \overline{u'v'} = -\kappa \frac{\partial^2 U}{\partial y^2} \quad \text{anti-diffusion!!}$$

Classical phenomenology attributes large-scale structure formation to turbulent cascades (inverse energy transfer from smaller to large scales)



We will deploy a statistical theory for the description of turbulent flows

Emergence of large-scale structure out of homogeneous turbulence will be understood as a statistical instability of the turbulent flow

Classical hydrodynamic stability



Lord Rayleigh

Lord Rayleigh taught us how to study the stability of a laminar flow to infinitesimal eddies

Can we study the stability of turbulent flows?

Classical hydrodynamic stability



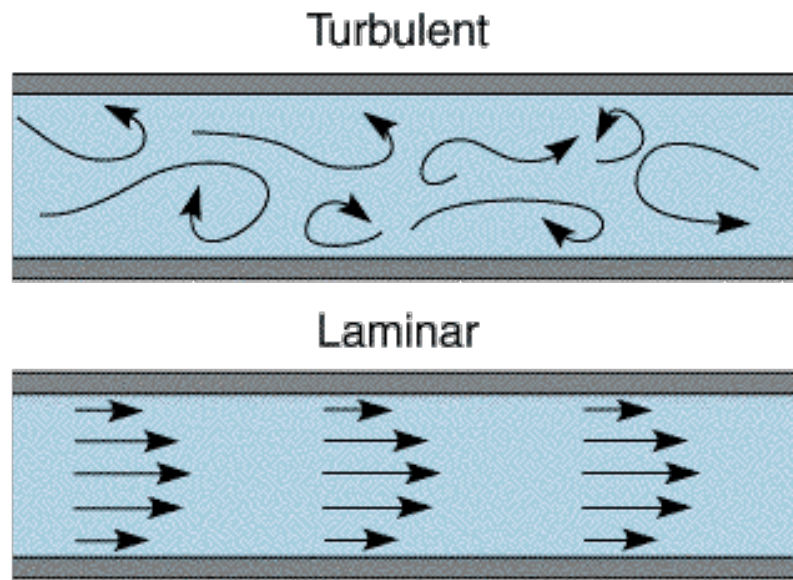
Lord Rayleigh

Hydrodynamic instabilities provide a way for eddies to gain energy from mean flow

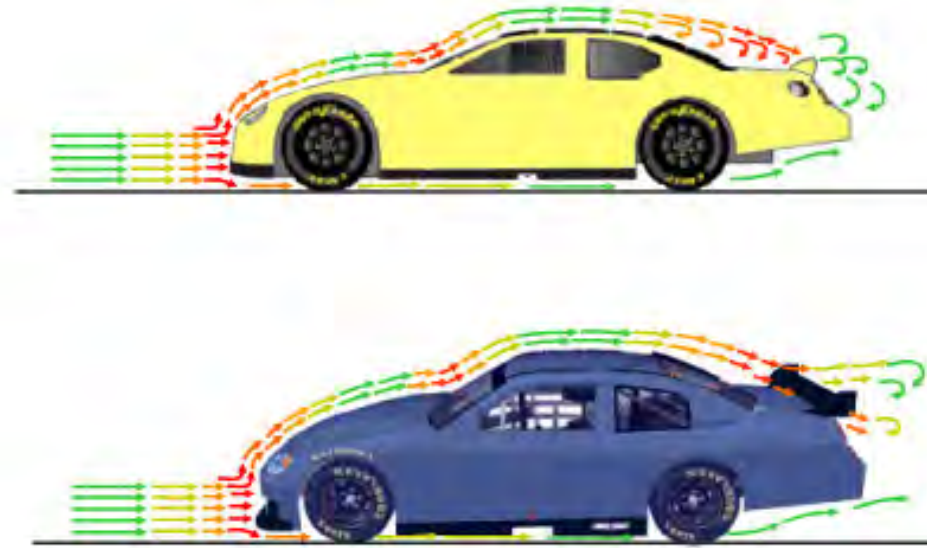
How about the opposite ?

Can the mean flow gain energy from the eddies through an instability ?

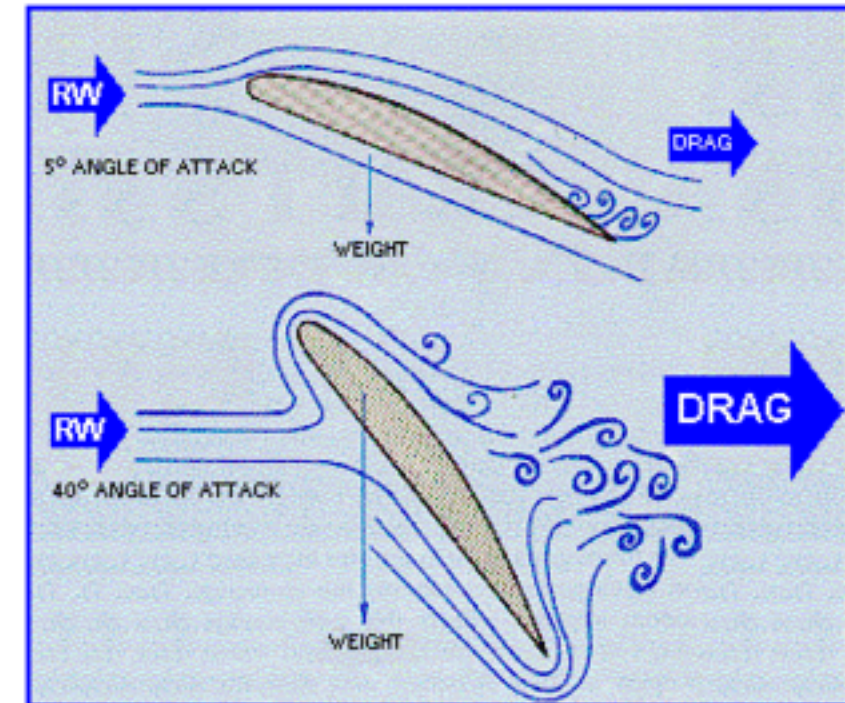
Turbulence (usually) acts as a drag



wall-bounded
flow



airflow over
vehicle



airflow over airfoil

can turbulence act to reinforce large scale flows?

Barotropic vorticity equation on a beta-plane

$$\partial_t \zeta + \mathbf{u} \cdot \nabla \zeta + \mathbf{u} \cdot \boldsymbol{\beta} = -r\zeta + \sqrt{\varepsilon} \xi$$

$$\nabla \cdot \mathbf{u} = 0$$

$$\nabla \times \mathbf{u} = \zeta \hat{\mathbf{z}}$$

$$\mathbf{u} = \hat{\mathbf{z}} \times \Delta^{-1} \zeta$$

$$(\Delta \equiv \nabla \cdot \nabla)$$

anisotropy
due to rotation

linear
dissipation
at rate r

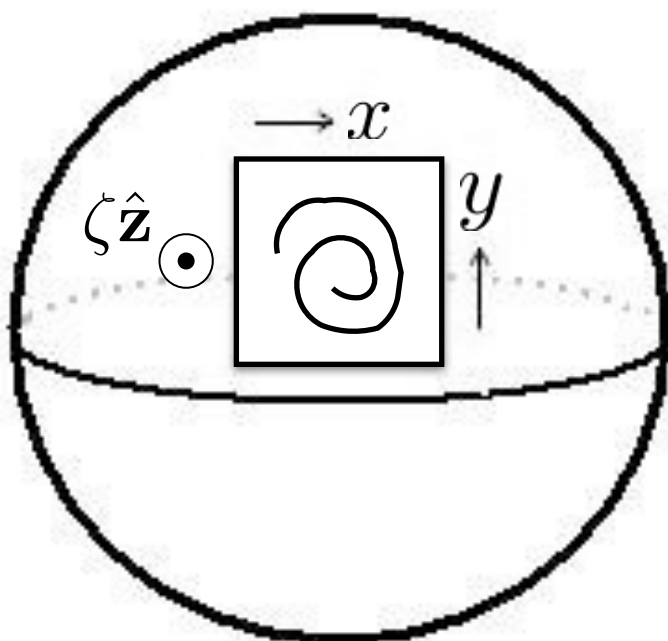
stochastic
forcing

$$\langle \xi(\mathbf{x}_a, t) \xi(\mathbf{x}_b, t') \rangle = Q(\mathbf{x}_a - \mathbf{x}_b) \delta(t - t')$$

ξ is statistically
homogeneous

$$\boldsymbol{\beta} = (0, \beta)$$

$\boldsymbol{\beta}$ is the gradient of
the planetary vorticity

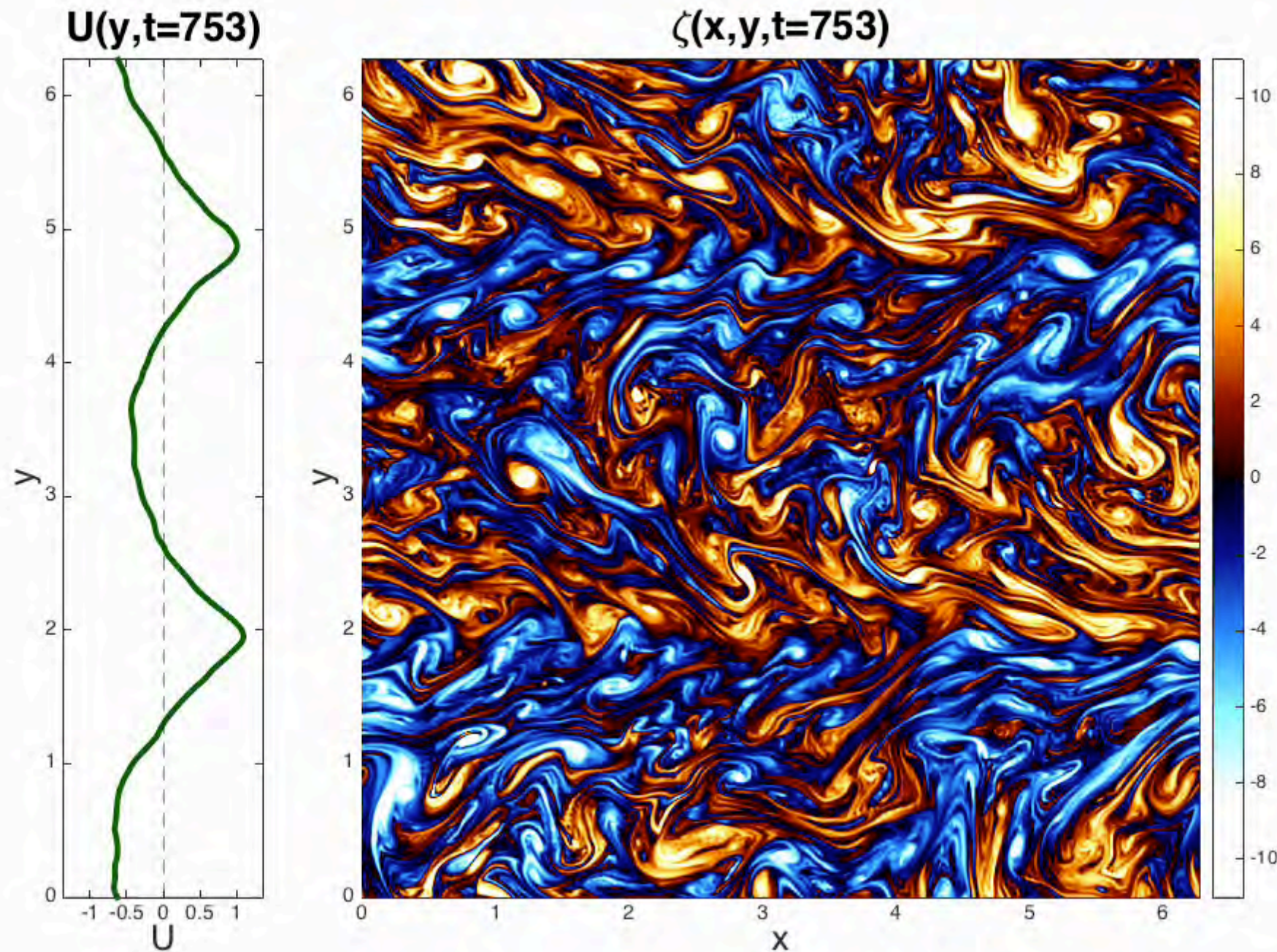


we have two non-
dimensional parameters

$$\varepsilon k_f^2 / r^3$$

$$\beta / (k_f r)$$

Barotropic β -plane turbulence exhibits large-scale structure formation



$$\varepsilon k_f^2 / r^3 = 10^6$$

$$\beta / (k_f r) = 67$$

statistically
homogeneous forcing

(no inhomogeneity is
imposed by the forcing)

initial random flow
inhomogeneities organize
the turbulence in a manner
so that they are reinforced

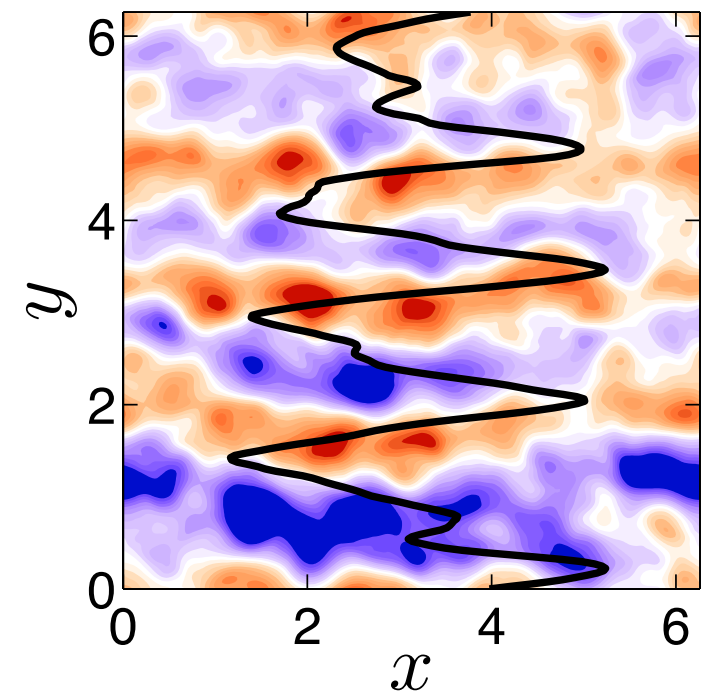
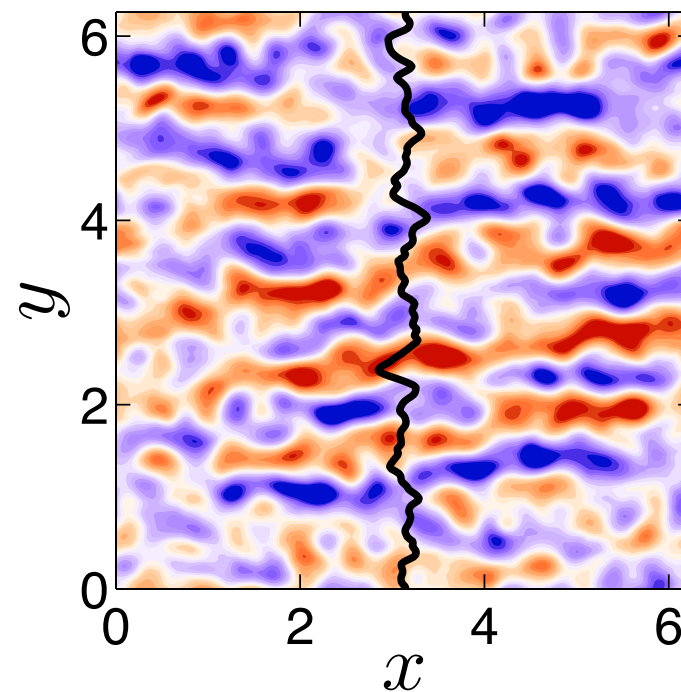
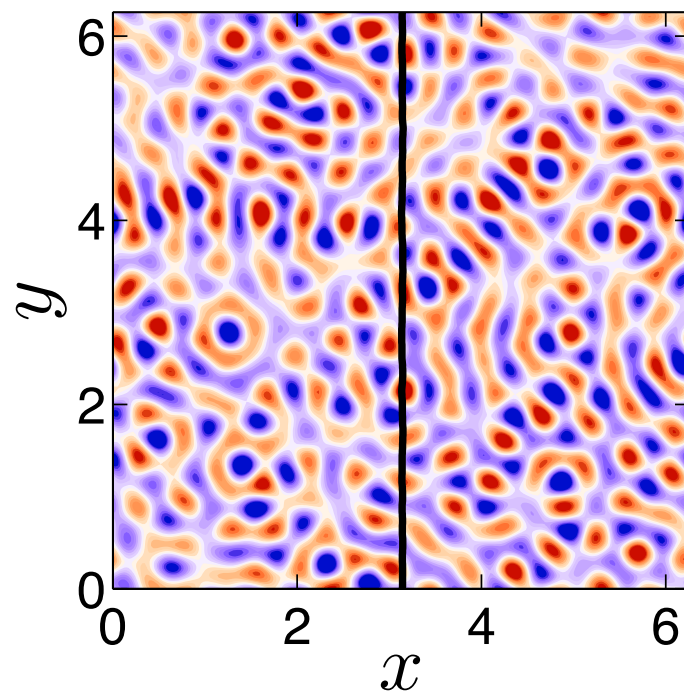
β -plane turbulence shows flows
at statistically steady state:
homogeneous — traveling waves — zonal jets

$$\beta/(k_f r) = 67$$

$$\varepsilon k_f^2 / r^3 = 10^2$$

$$5 \times 10^3$$

$$5 \times 10^4$$



this suggests that there is some kind of transition as ε is increased

[shown are snapshots of the streamfunction field $\psi(\mathbf{x}, t)$
with instantaneous zonal mean flow $U(y, t)$]

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Barotropic vorticity equation on a beta-plane

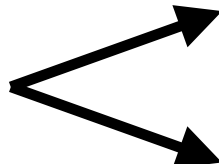
Using decomposition: $\zeta(\mathbf{x}, t) = \underbrace{\mathcal{T}[\zeta(\mathbf{x}, t)]}_{\substack{Z(\mathbf{x}, t) \\ \text{mean flow}}} + \underbrace{\zeta'(\mathbf{x}, t)}_{\text{eddies}}$

$$\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} = -\mathcal{T}[\mathbf{u}' \cdot \nabla \zeta'] - rZ$$

$$\partial_t \zeta' = \mathcal{A}(\mathbf{U}) \zeta' + \mathcal{T}[\mathbf{u}' \cdot \nabla \zeta'] - \mathbf{u}' \cdot \nabla \zeta' + \sqrt{\varepsilon} \xi$$

with

$$\mathcal{A}(\mathbf{U}) \equiv -\mathbf{U} \cdot \nabla + \left[(\Delta \mathbf{U}) - \boldsymbol{\beta} \cdot (\hat{\mathbf{z}} \times \nabla) \right] \Delta^{-1} - r$$

\mathcal{T}  average over the zonal direction x
 Reynolds over an intermediate time scale or length scale
 (larger than the time scale or length scale of the turbulent motions
 and smaller than the time scale or length scale of mean field)

NL system

$$\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} = -\mathcal{T} [\mathbf{u}' \cdot \nabla \zeta'] - rZ$$

$$\partial_t \zeta' = \mathcal{A}(\mathbf{U}) \zeta' + \mathcal{T} [\mathbf{u}' \cdot \nabla \zeta'] - \mathbf{u}' \cdot \nabla \zeta' + \sqrt{\varepsilon} \xi$$

NL system

restrict nonlinearity by *not* allowing
eddy-eddy \rightarrow eddy interactions (QL)

$$\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} = -\mathcal{T} [\mathbf{u}' \cdot \nabla \zeta'] - rZ$$

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QL system

restrict nonlinearity by *not* allowing
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$$\partial_t \zeta' = \mathcal{A}(\mathbf{U}) \zeta' + \sqrt{\varepsilon} \xi$$

QL allows **only** the direct, two-way interaction
of the eddies and the mean flow

QL does **not** include turbulent cascades

S3T system

Under the ergodic assumption that the average T is equal to ensemble average over forcing realizations:

$$\mathcal{T}(\bullet) = \langle \bullet \rangle$$

we derive from QL a *closed* system for the evolution of the first two statistical moments of the flow

$$Z(\mathbf{x}, t) = \langle \zeta(\mathbf{x}, t) \rangle \quad , \quad C(\mathbf{x}_a, \mathbf{x}_b, t) = \langle \zeta'(\mathbf{x}_a, t) \zeta'(\mathbf{x}_b, t) \rangle$$

S3T system

$$\begin{aligned}\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} &= \mathcal{R}(C) - rZ \\ \partial_t C_{ab} &= [\mathcal{A}_a(\mathbf{U}) + \mathcal{A}_b(\mathbf{U})] C_{ab} + \varepsilon Q_{ab}\end{aligned}$$

with

$$C_{ab} \equiv C(\mathbf{x}_a, \mathbf{x}_b, t) = \langle \zeta'(\mathbf{x}_a, t) \zeta'(\mathbf{x}_b, t) \rangle$$

$$Q_{ab} \equiv Q(\mathbf{x}_a - \mathbf{x}_b) \longrightarrow \text{the spatial covariance of the statistically homogeneous stochastic forcing}$$

$$\mathcal{R}(C) \equiv -\langle \mathbf{u}' \cdot \nabla \zeta' \rangle = -\nabla \cdot \left[\frac{\hat{\mathbf{z}}}{2} \times (\nabla_a \Delta_a^{-1} + \nabla_b \Delta_b^{-1}) C_{ab} \right]_{a=b}$$

(the Reynolds stresses are given as a linear function of C)

S3T system

$$\begin{aligned}\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} &= \mathcal{R}(C) - rZ \\ \partial_t C_{ab} &= [\mathcal{A}_a(\mathbf{U}) + \mathcal{A}_b(\mathbf{U})] C_{ab} + \varepsilon Q_{ab}\end{aligned}$$

Neglect of the eddy-eddy term in NL is equivalent with neglect of third and higher-order statistical moments.

S3T system

(the theory)

$$\begin{aligned}\partial_t Z + \mathbf{U} \cdot \nabla Z + \mathbf{U} \cdot \boldsymbol{\beta} &= \mathcal{R}(C) - rZ \\ \partial_t C_{ab} &= [\mathcal{A}_a(\mathbf{U}) + \mathcal{A}_b(\mathbf{U})] C_{ab} + \varepsilon Q_{ab}\end{aligned}$$

The S3T system

- autonomous
- deterministic
- admits fixed point solutions consisting of a mean flow and second-order eddy statistics $(\mathbf{U}^e(\mathbf{x}), C^e(\mathbf{x}_a, \mathbf{x}_b))$
- allows the study of the stability of such equilibrium solutions

stability of S3T equilibria

$$\left(\mathbf{U}^e(\mathbf{x}), C^e(\mathbf{x}_a, \mathbf{x}_b) \right)$$

perturbations $(\delta Z, \delta C)$ about an equilibrium satisfy the linearized S3T equations:

$$\begin{aligned} \partial_t \delta Z &= \mathcal{A}^e \delta Z + \mathcal{R}(\delta C) \\ \partial_t \delta C_{ab} &= (\mathcal{A}_a^e + \mathcal{A}_b^e) \delta C_{ab} + (\delta \mathcal{A}_a + \delta \mathcal{A}_b) C_{ab}^e \end{aligned} \quad \mathcal{A}^e \equiv \mathcal{A}(\mathbf{U}^e)$$

eigenanalysis of this system determines the stability of $\left(\mathbf{U}^e(\mathbf{x}), C^e(\mathbf{x}_a, \mathbf{x}_b) \right)$

stability of S3T equilibria

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hydrodynamic
stability

$$\partial_t \delta Z = \mathcal{A}^e \delta Z + \mathcal{R}(\delta C)$$

$$\partial_t \delta C_{ab} = (\mathcal{A}_a^e + \mathcal{A}_b^e) \delta C_{ab} + (\delta \mathcal{A}_a + \delta \mathcal{A}_b) C_{ab}^e$$

$$\mathcal{A}^e \equiv \mathcal{A}(\mathbf{U}^e)$$

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for statistically homogeneous forcing there exists *always*
a statistically homogeneous S3T equilibrium

$$\mathbf{U}^e = 0 \quad , \quad C^e(\mathbf{x}_a - \mathbf{x}_b) = \frac{\varepsilon Q}{2r} \quad \text{(for any } \varepsilon, \beta \text{ and homogeneous } Q)$$

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eigenfunctions:

$$\delta Z = e^{i\mathbf{n} \cdot \mathbf{x}} e^{(\sigma - i\omega_{\mathbf{n}})t} \quad \text{(plane wave)}$$

$$\delta C_{ab} = e^{i\mathbf{n} \cdot (\mathbf{x}_a + \mathbf{x}_b)/2} \delta C^{(h)}(\mathbf{x}_a - \mathbf{x}_b) e^{(\sigma - i\omega_{\mathbf{n}})t}$$

Instability occurs when $\text{Re}(\sigma) > 0$ for at least one plane wave \mathbf{n}

Eigenvalues σ satisfy:

$$\sigma + r = \varepsilon \int \frac{d^2\mathbf{k}}{(2\pi)^2} \frac{|\mathbf{n} \times \mathbf{k}|^2 (k_p^2 - k^2)(k^2 - n^2)}{k_p^2 k^4 n^2 [\sigma + 2r + i(\omega_{\mathbf{k}+\mathbf{n}} - \omega_{\mathbf{k}} - \omega_{\mathbf{n}})]} \frac{\hat{Q}(\mathbf{k})}{2r}$$

$$\mathbf{k}_p = \mathbf{k} + \mathbf{n} \quad , \quad k_p = |\mathbf{k}_p| \quad , \quad k = |\mathbf{k}| \quad , \quad \omega_{\mathbf{n}} = \frac{-\beta n_x}{n^2} \quad \text{Rossby wave frequency} \quad , \quad \hat{Q}(\mathbf{k}) = \int \frac{d^2\mathbf{k}}{(2\pi)^2} Q(\mathbf{x}_a - \mathbf{x}_b) e^{i\mathbf{k} \cdot (\mathbf{x}_a - \mathbf{x}_b)}$$

take forcing prescribed
with spatial covariance with spectrum

$$\hat{Q}(\mathbf{k}) \sim \delta(k - k_f) \left[1 + \mu \frac{k_x^2 - k_y^2}{k^2} \right] \quad |\mu| \leq 1$$

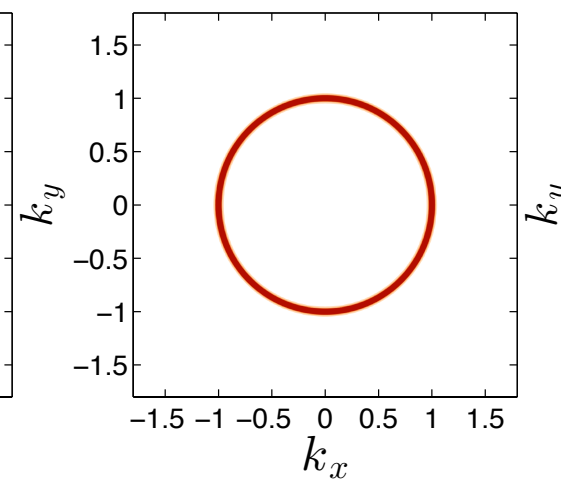
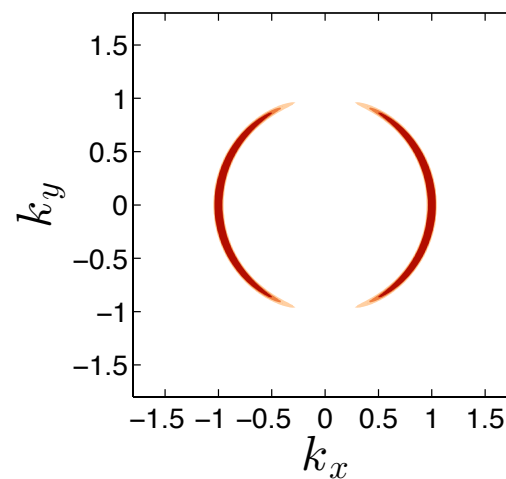
anisotropic
(baroclinic)
[\approx Earth]

isotropic
(convection)
[\approx Jupiter]

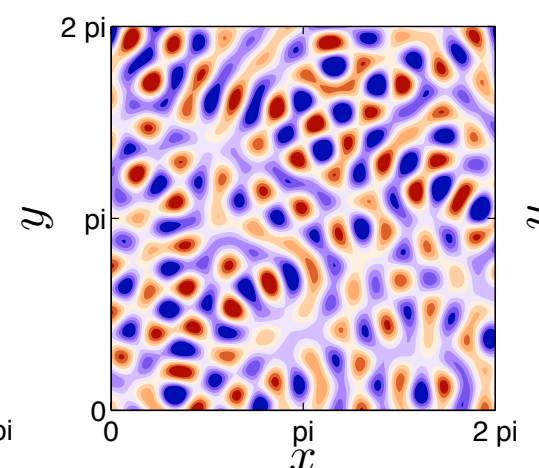
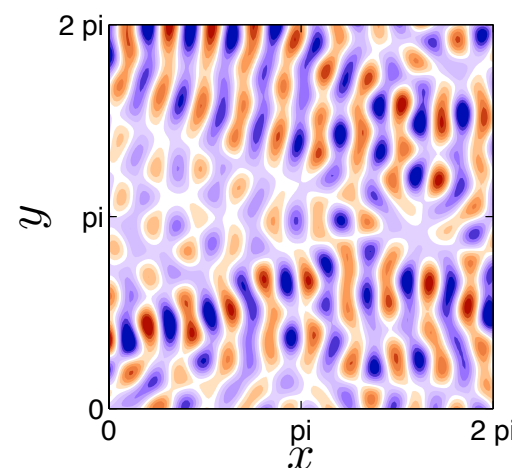
$$\mu = 1$$

$$\mu = 0$$

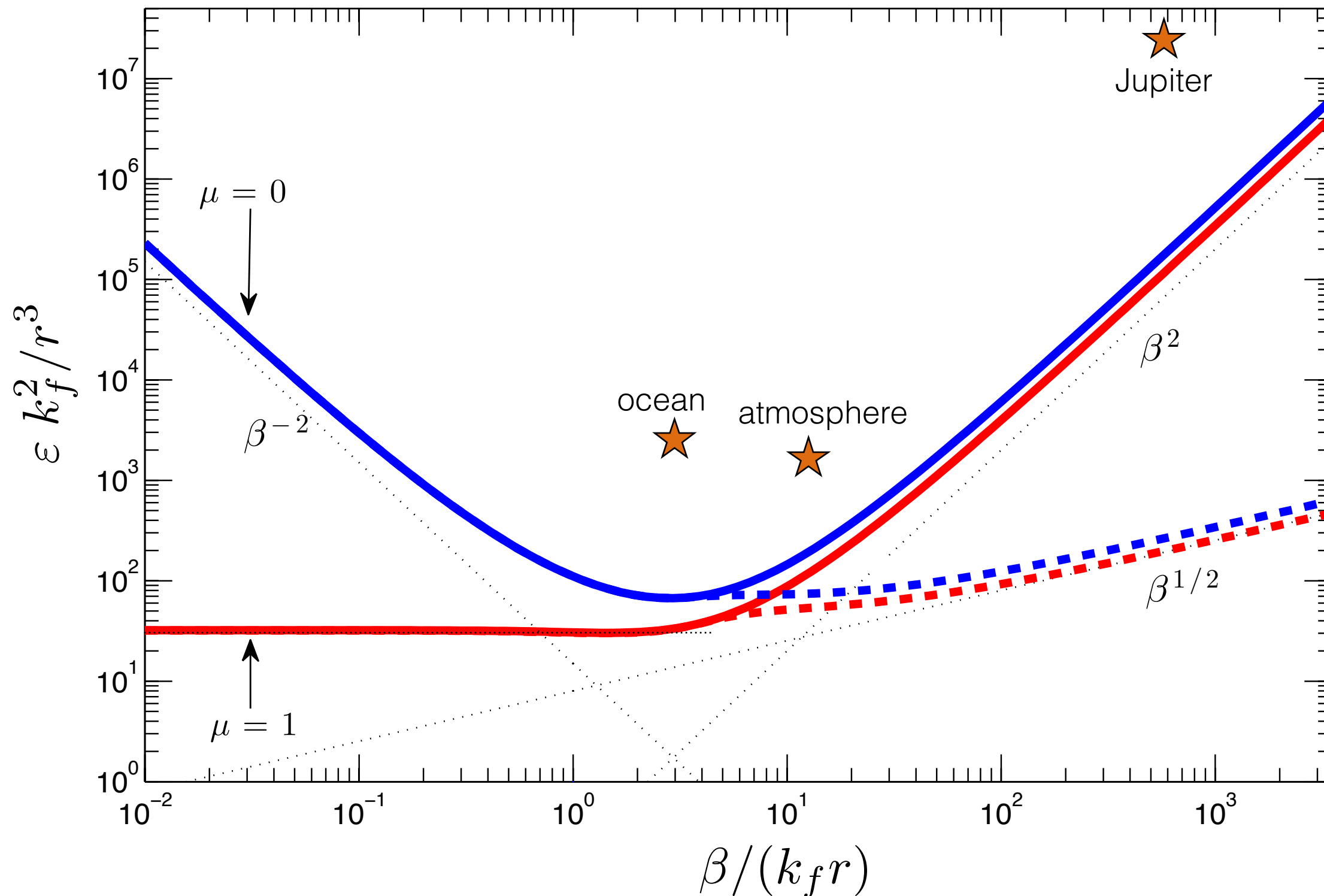
$\hat{Q}(\mathbf{k})$



$\xi(\mathbf{x}, t)$

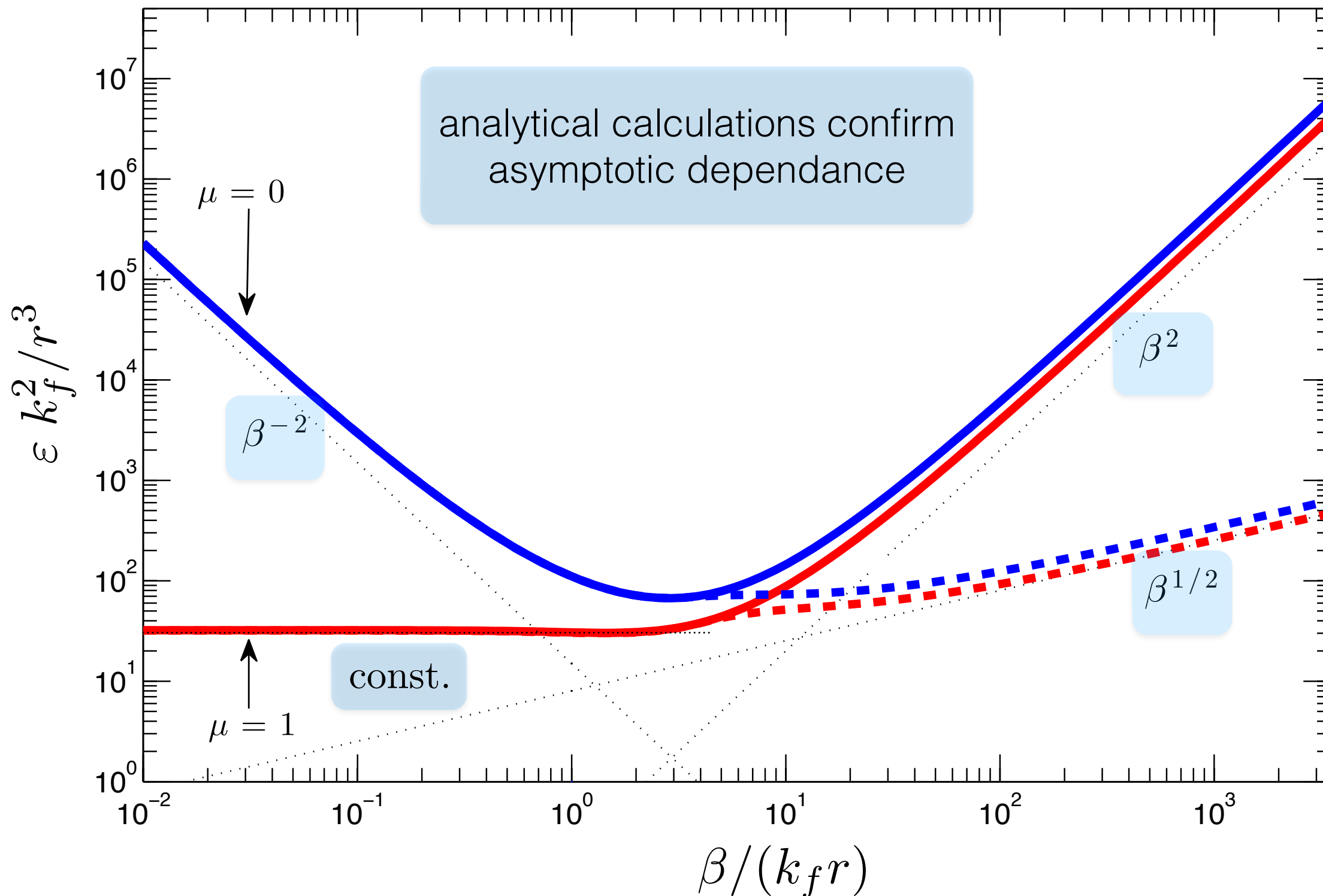


Critical ε for S3T instability of the homogeneous state



for large β the critical ε scaling is independent of the forcing structure

Critical ε for S3T instability of the homogeneous state



for large β the critical ε scaling is independent of the forcing structure

instability of the homogeneous turbulent state

$$\sigma + r = \underbrace{\varepsilon \int \frac{d^2 \mathbf{k}}{(2\pi)^2} \frac{|\mathbf{n} \times \mathbf{k}|^2 (k_p^2 - k^2)(k^2 - n^2)}{k_p^2 k^4 n^2 [\sigma + 2r + i(\omega_{\mathbf{k}+\mathbf{n}} - \omega_{\mathbf{k}} - \omega_{\mathbf{n}})]} \frac{\hat{Q}(\mathbf{k})}{2r}}_f$$

mean flow perturbation equation

$$(\partial_t - \mathcal{A}^e) \delta Z = \mathcal{R}(\delta C)$$

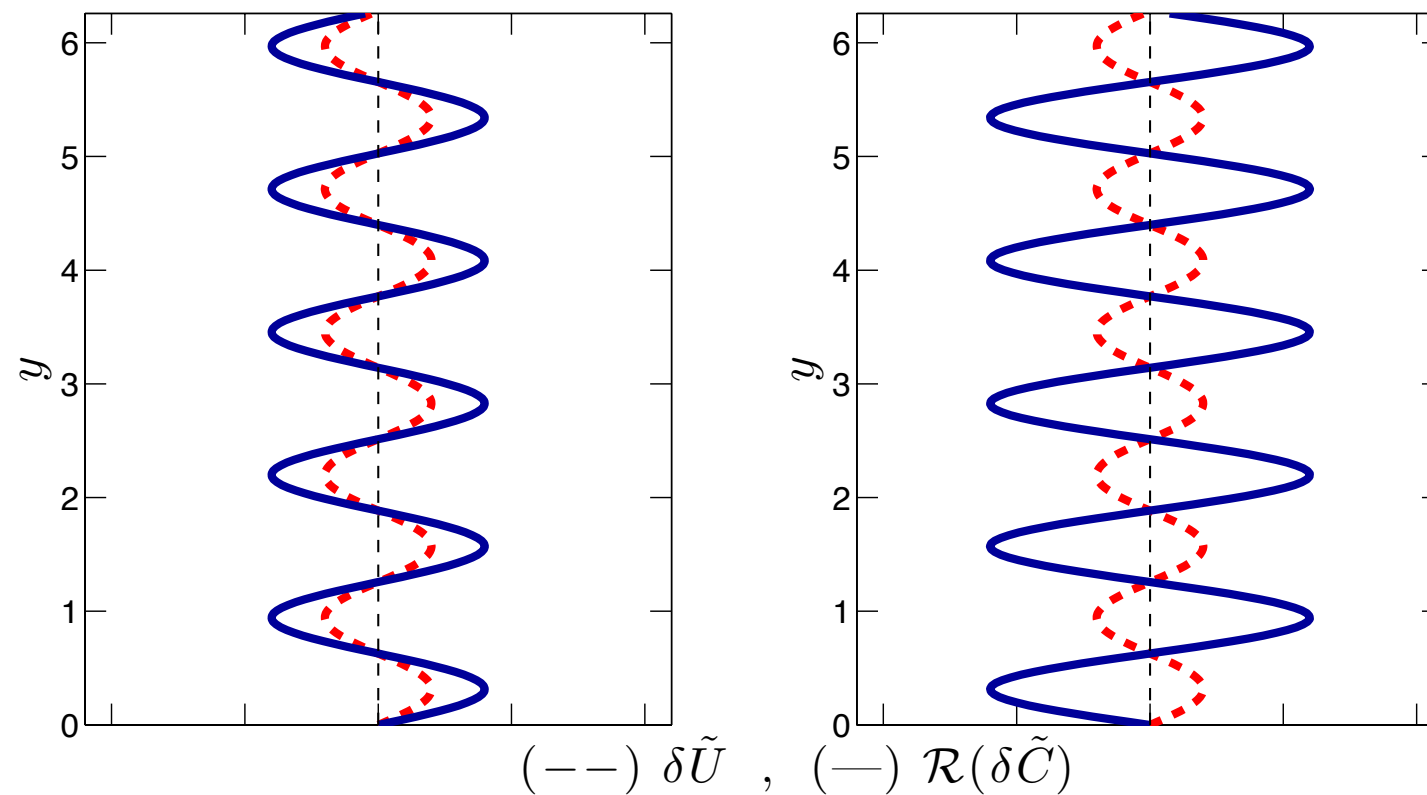
f is the sum of the contributions of the spectral components of Q (or C^e) to the perturbation Reynolds stress divergence

$$\text{Instability requires} \quad f_r \equiv \text{Re}(f) \Big|_{\varepsilon_c} > 0 \quad \varepsilon_c = \frac{r}{f_r}$$

f_r determines if turbulence will act so as to reinforce or diminish the infinitesimal mean flow

$$f_r > 0$$

$$f_r < 0$$



turbulence acts as:

anti-diffusion

diffusion

analytic calculations for f_r show that
turbulence acts anti-diffusively
even for infinitesimal mean flows!

small β

anisotropic forcing ($\mu \neq 0$) $\longrightarrow f_r \sim \mu \longrightarrow \varepsilon_c \sim 1/\mu$

isotropic forcing ($\mu = 0$) $\longrightarrow f_r \sim \beta^2 \longrightarrow \varepsilon_c \sim \beta^{-2}$

$$(\partial_t + r)\delta U = -\partial_y \delta \langle u'v' \rangle = -C \mu \partial_{yy}^2 \delta U + D \beta^2 \partial_{yyyy}^4 \delta U \quad C, D > 0$$

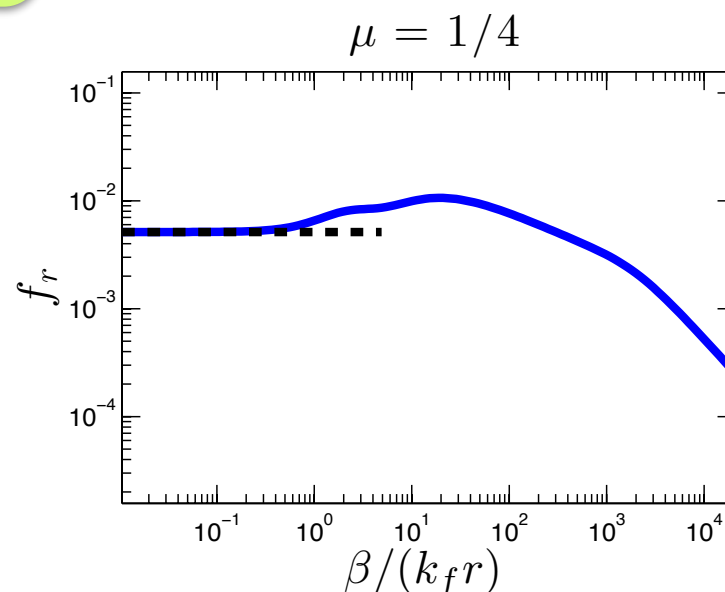
Jupiter's
finite amplitude jets

$$\partial_y \overline{u'v'} = \kappa \partial_{yy}^2 U$$

$$\kappa \approx 10^6 \text{ m}^2 \text{ s}^{-1}$$

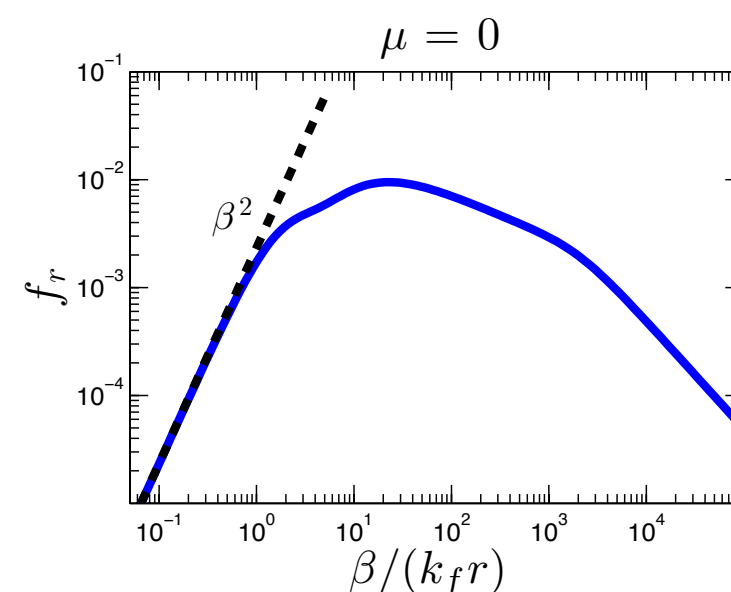
$\mu > 0$

2nd order
anti-diffusion



$\mu = 0$

4th order
hyper-anti-diffusion

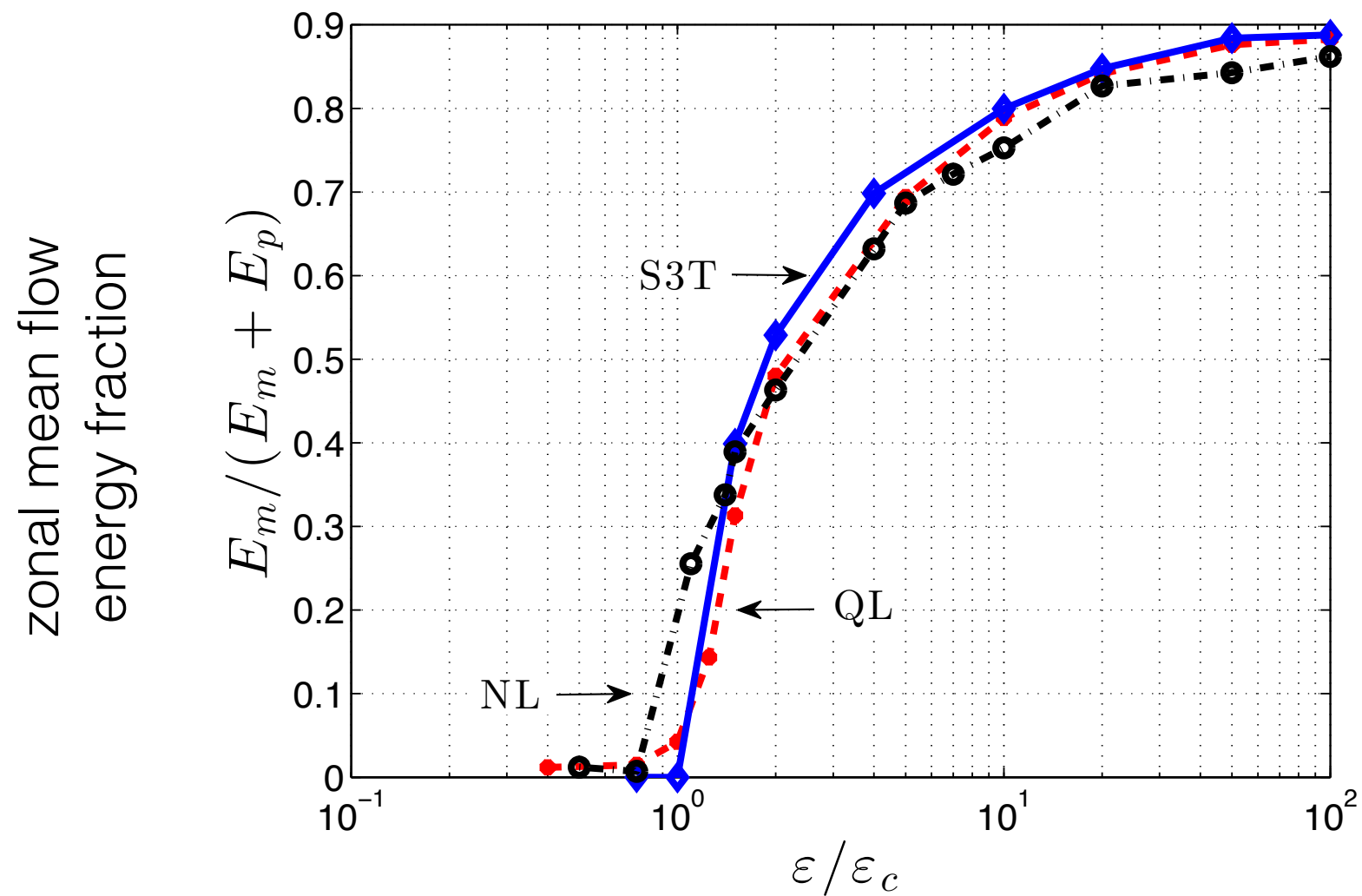


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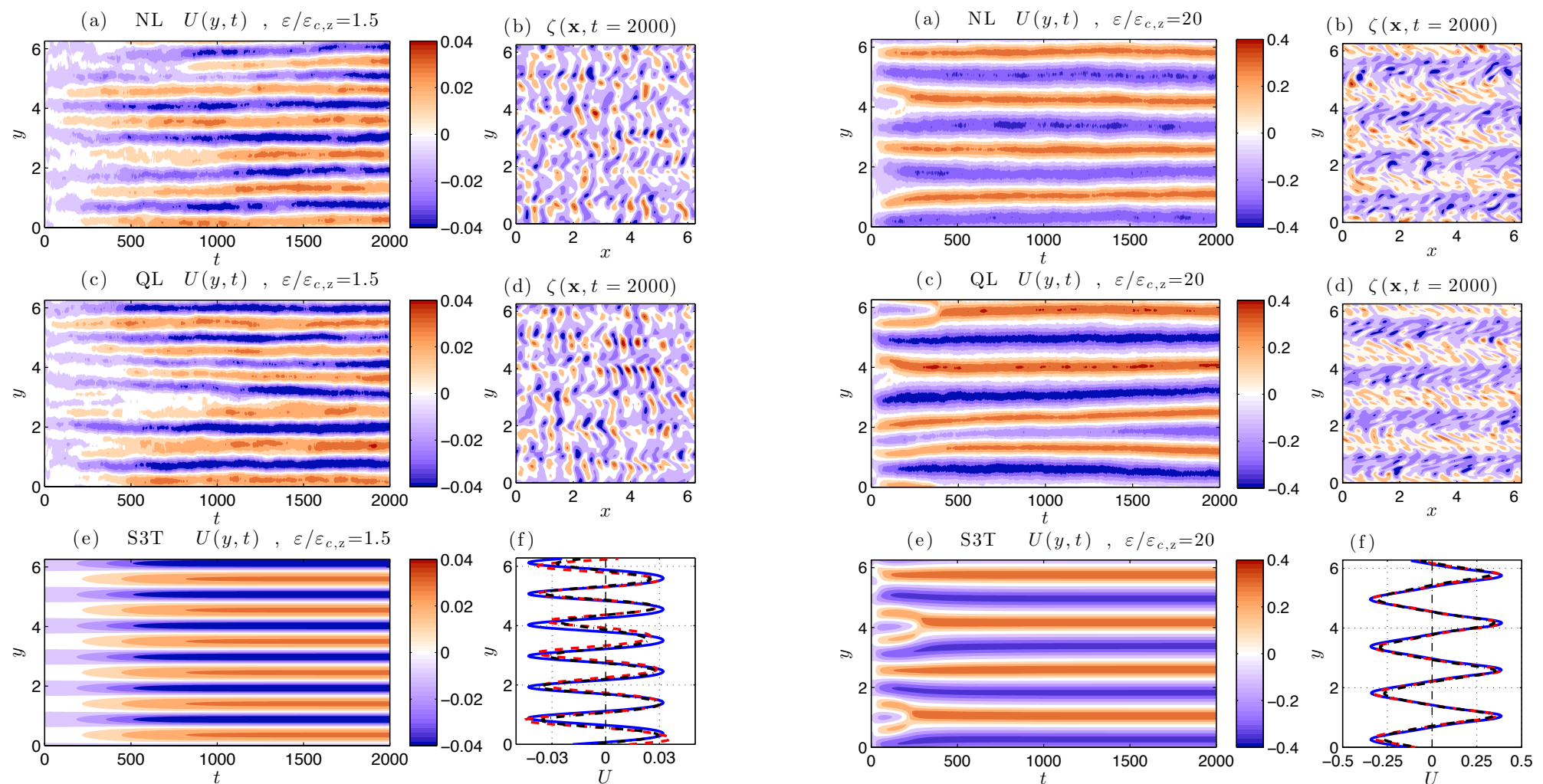
S3T predictions for jet formation and equilibration at finite amplitude

anisotropic forcing
[Earth-like]



S3T predictions for jet formation and equilibration at finite amplitude

$\varepsilon/\varepsilon_{c,z} = 1.5$ anisotropic forcing [Earth-like] $\varepsilon/\varepsilon_{c,z} = 20$



statistical instabilities that are predicted by S3T show up in single NL/QL realizations of the flow

emergent instabilities grow and reach finite amplitude

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Zonal jet S3T equilibria

$$\mathbf{U}^e(\mathbf{x}) = \left(U^e(y), 0 \right) \quad , \quad C^e(x_a - x_b, y_a, y_b)$$

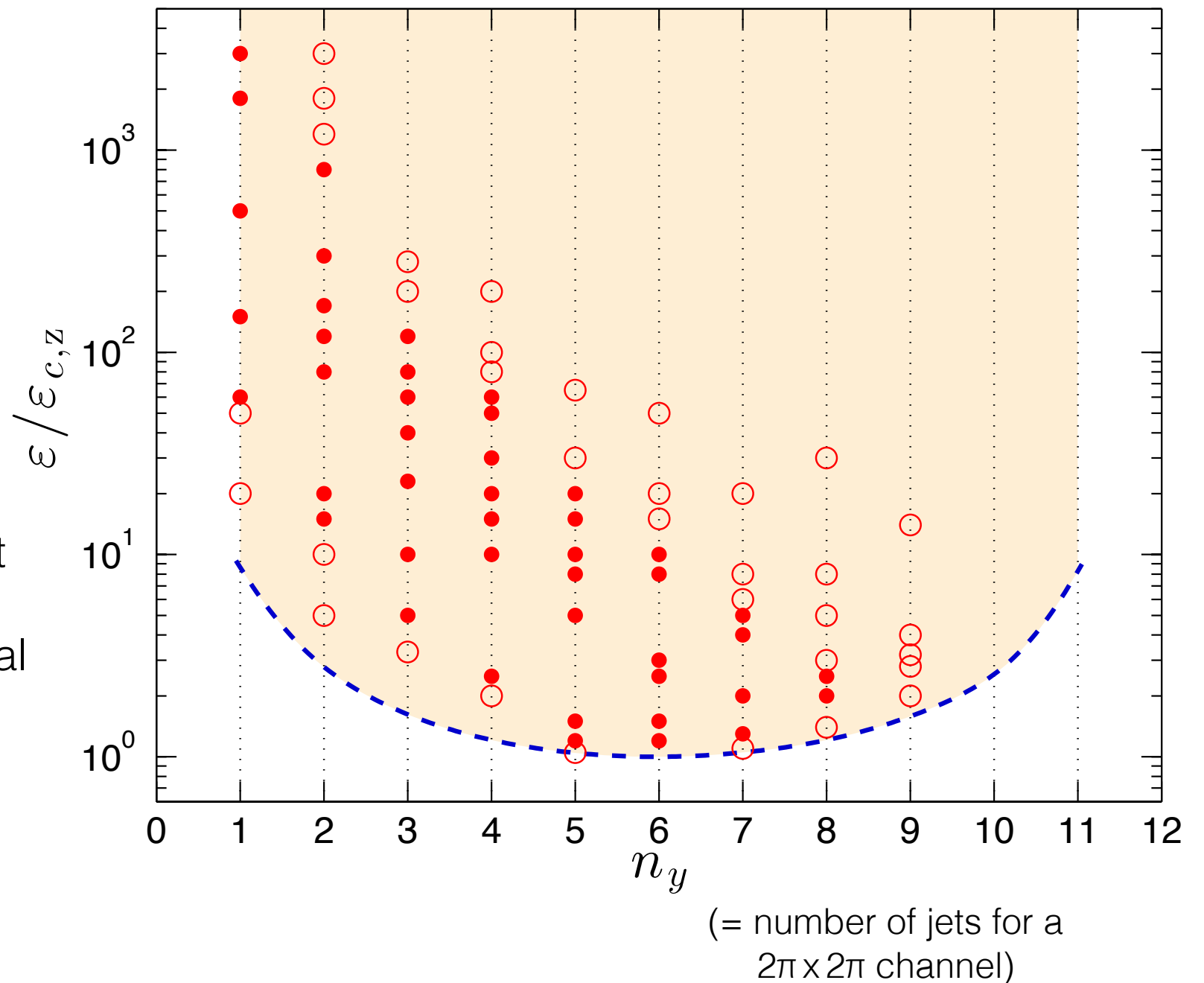
We have developed numerical methods for

- i) determining such equilibria with great accuracy and
- ii) studying their S3T stability

Stability of zonal jet S3T equilibria to zonal jet perturbations

Stability analysis of
inhomogeneous turbulent
states with zonal jets predicts:

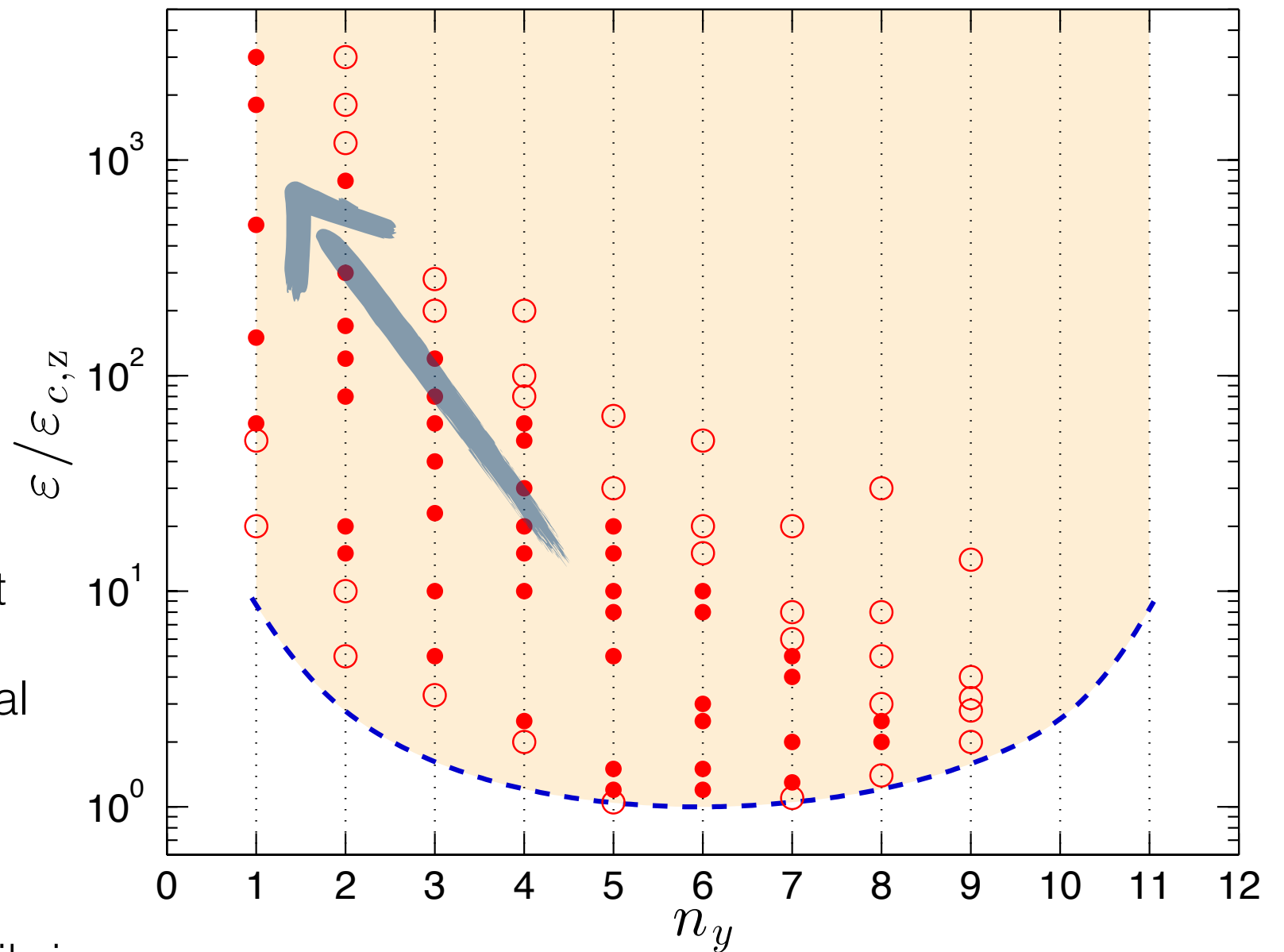
- ▶ existence of multiple equilibria
and their domain of attraction
- ▶ merging of jets as ε increases
- ▶ finite amplitude equilibration at
small supercriticality is
described through the universal
Eckhaus instability of the G-L
amplitude equation



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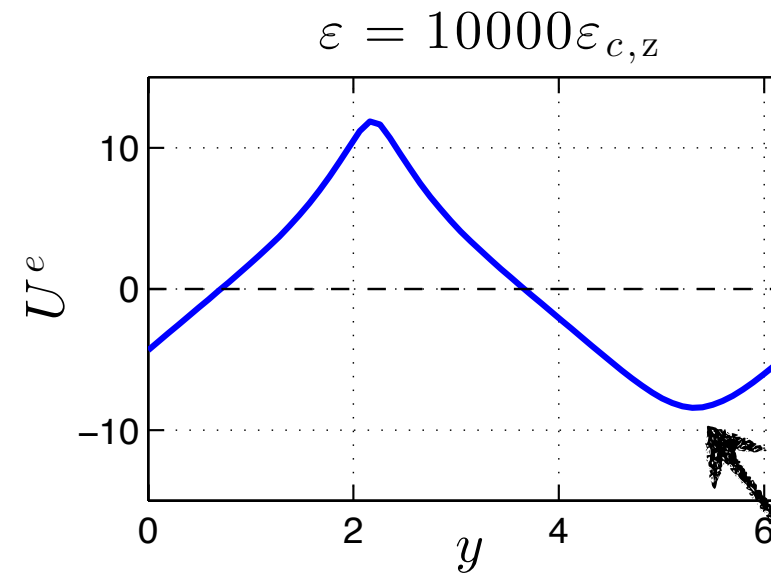
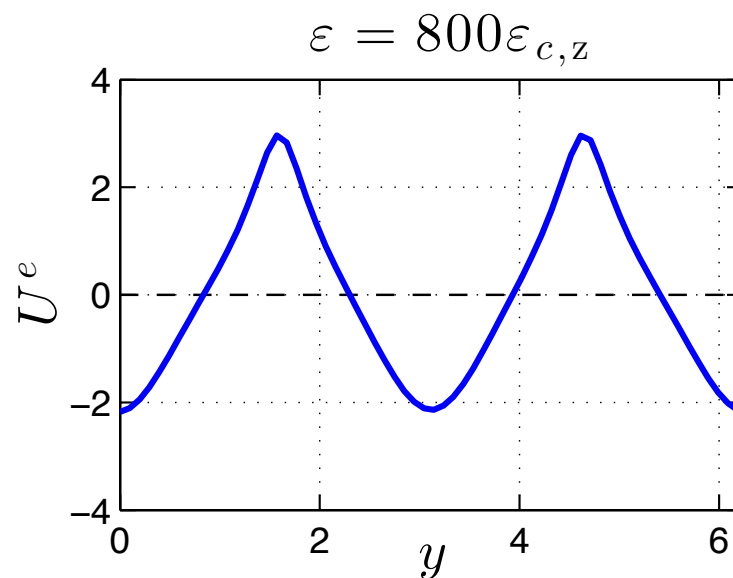
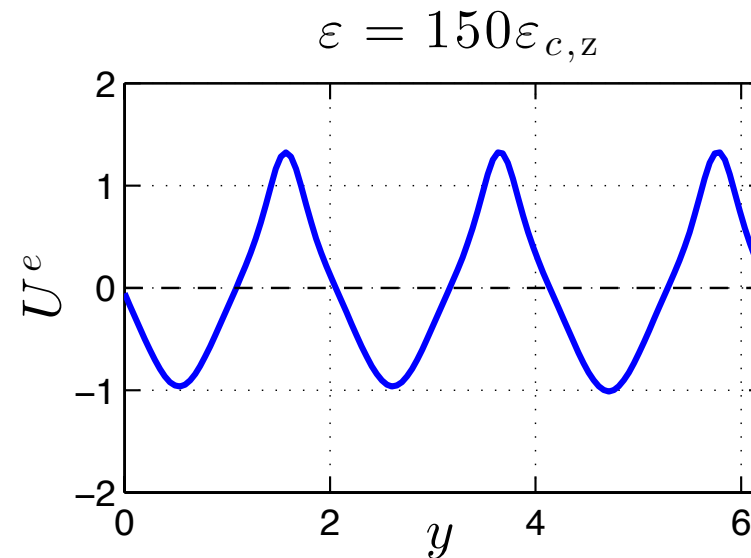
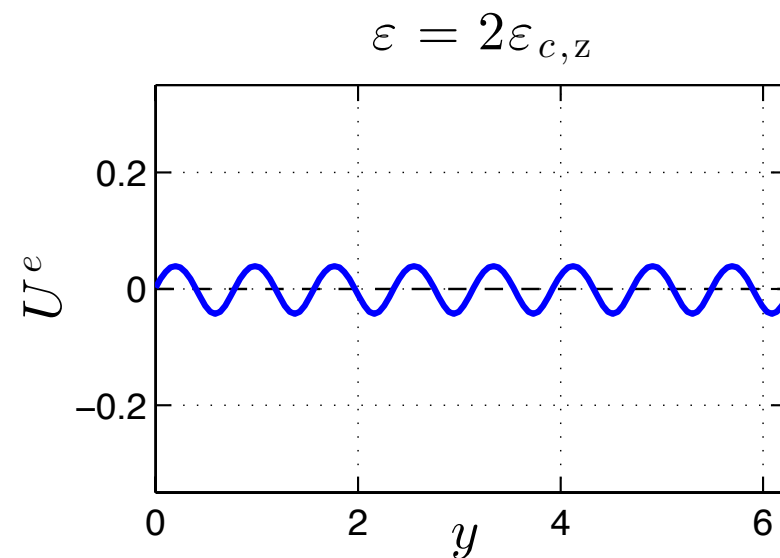
- ▶ existence of multiple equilibria and their domain of attraction
- ▶ merging of jets as ε increases
- ▶ finite amplitude equilibration at small supercriticality is described through the universal Eckhaus instability of the G-L amplitude equation



For higher energy input rates equilibria become S3T unstable and move towards the left of the diagram

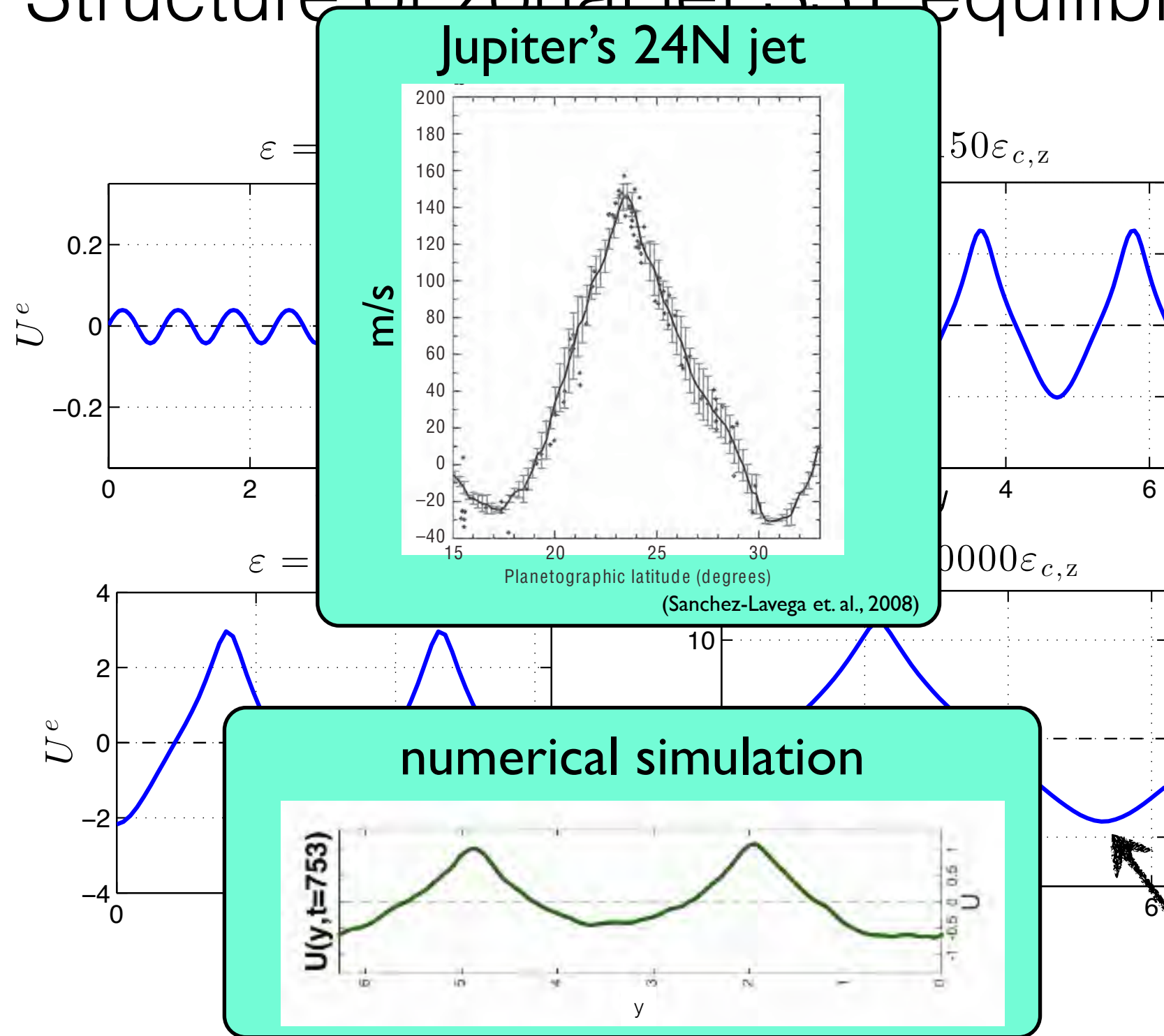
(= number of jets for a $2\pi \times 2\pi$ channel)

Structure of zonal jet S3T equilibria

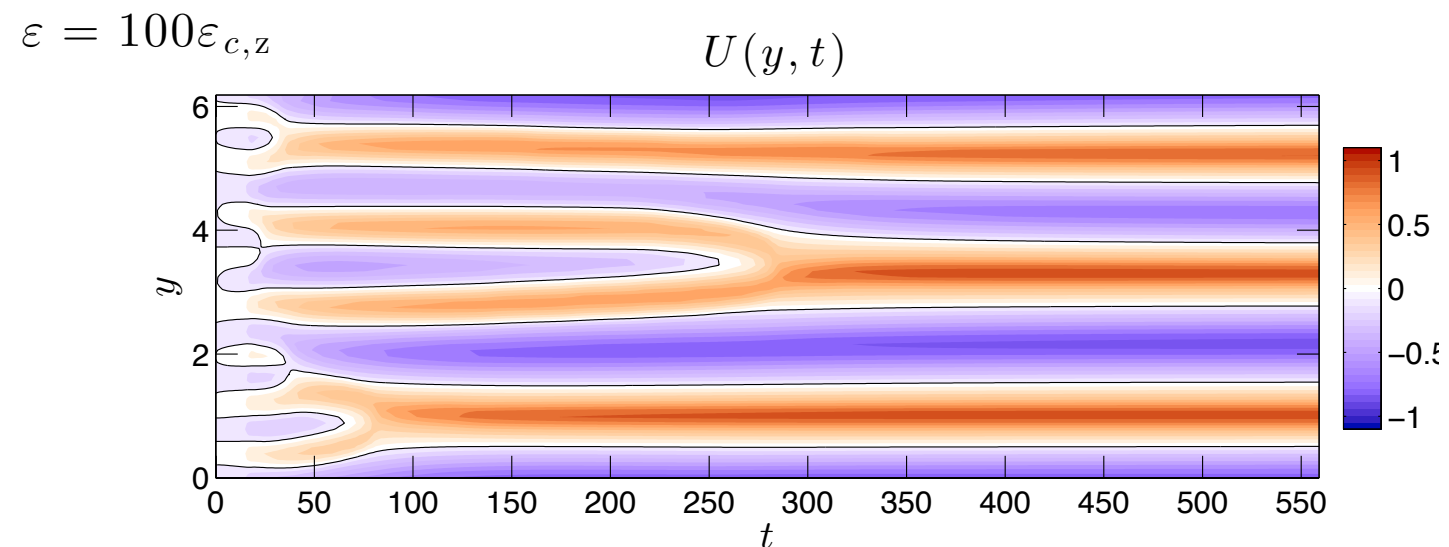


$\beta - d^2U/dy^2 \approx 0$
Rayleigh-Kuo hydrodynamic
stability criterion

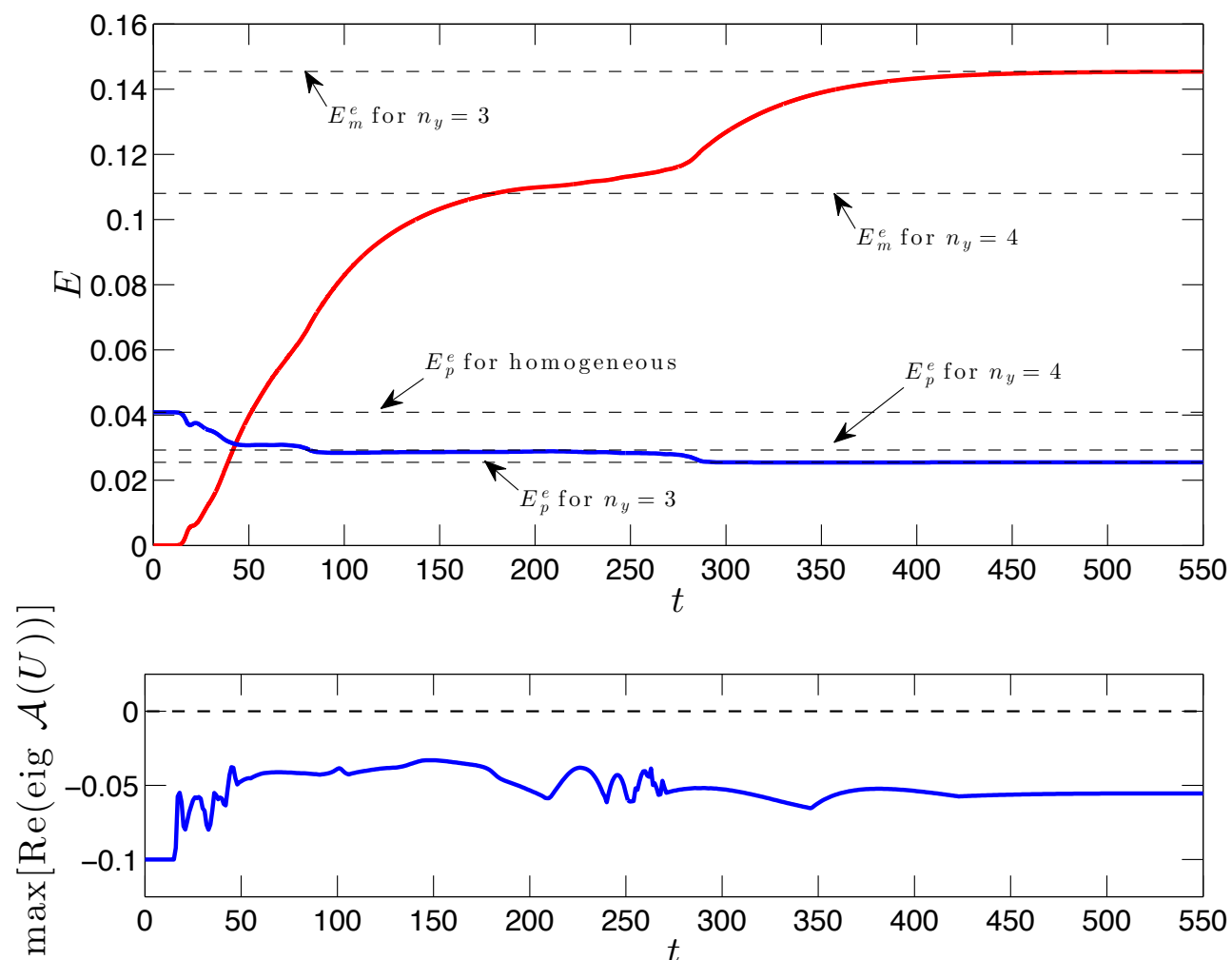
Structure of zonal jet S3T equilibria



Jet mergers do *not* occur due to hydrodynamic instabilities

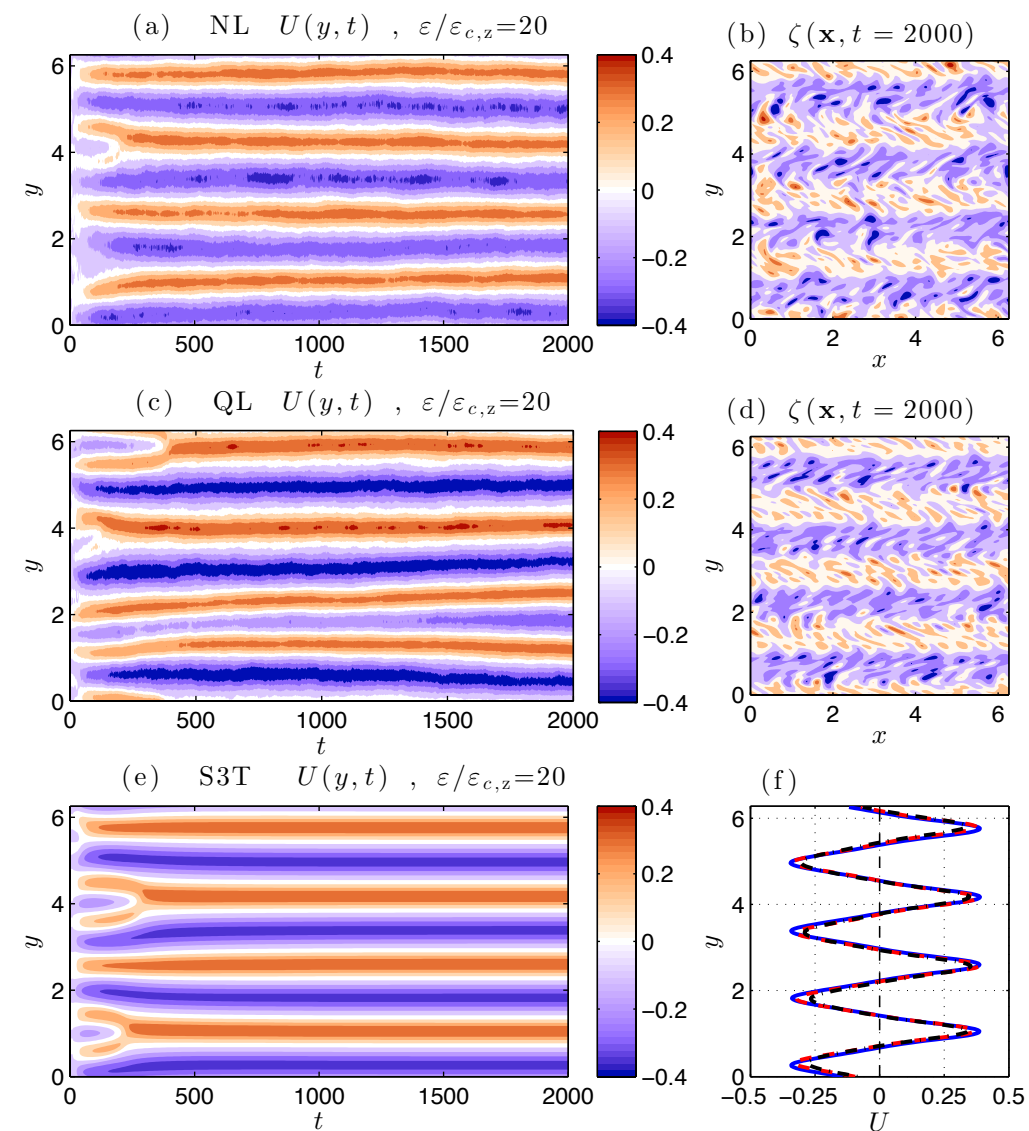


at this supercriticality
the $n_y=4$ jet equilibrium
is S3T unstable



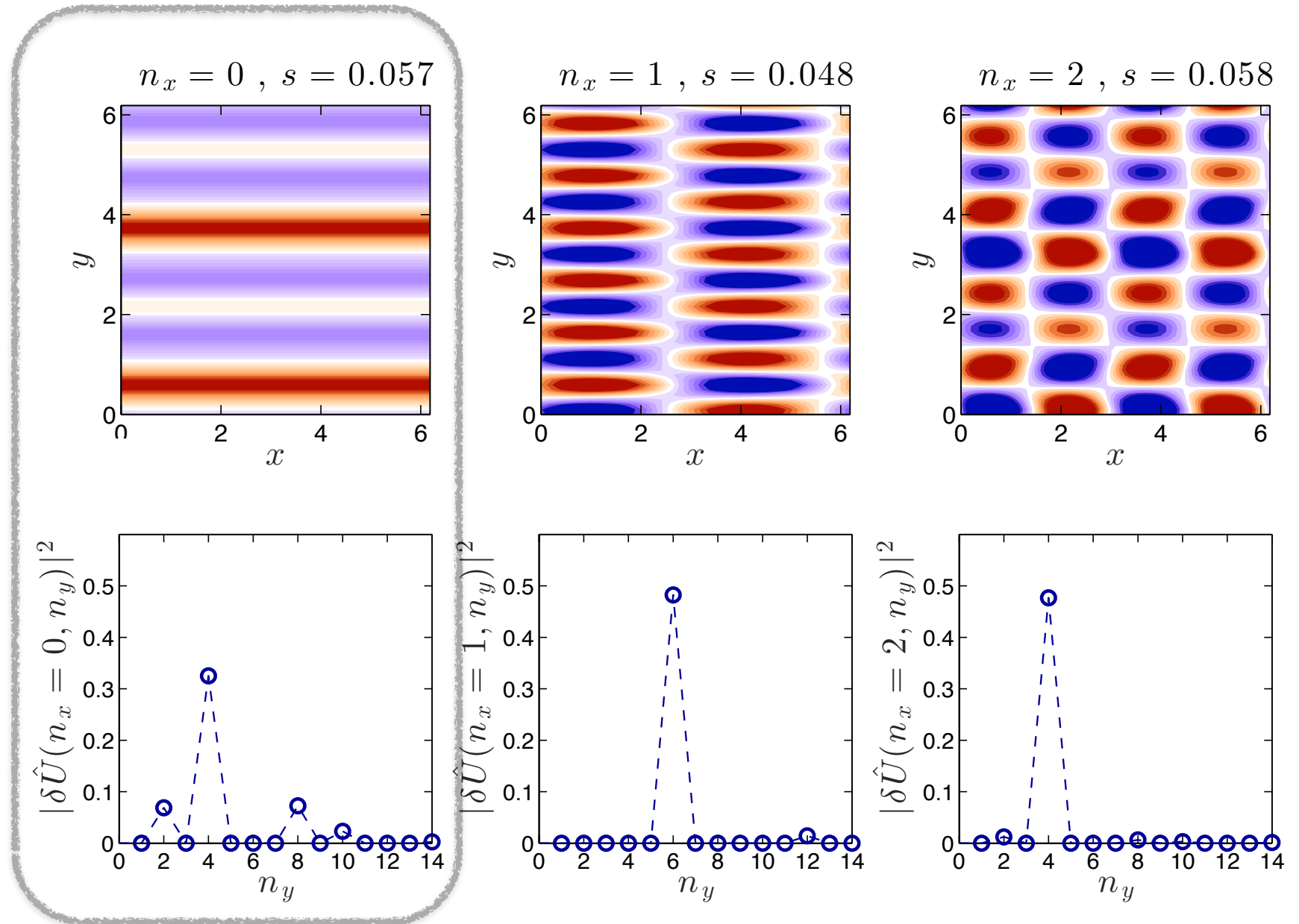
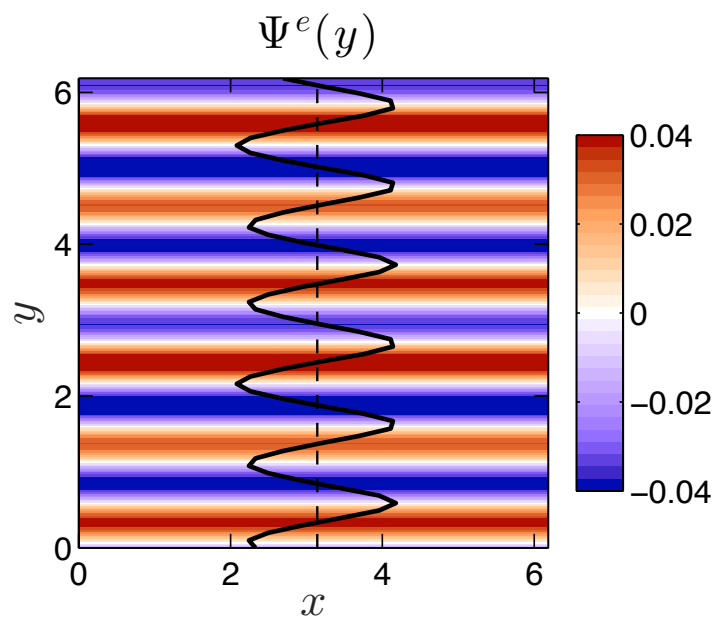
S3T stability analysis of the jet equilibria predicts the jet merger as the most unstable eigenfunction

Remember the case with $\varepsilon/\varepsilon_{c,z} = 20$



S3T stability analysis of the jet equilibria predicts the jet merger as the most unstable eigenfunction

$n_y = 6$ jet equilibrium at
 $\varepsilon/\varepsilon_{c,z} = 20$



structure of talk

- ▶ introduction to the physical problem
- ▶ formulation of the theory (S3T)
- ▶ study of the stability of homogeneous turbulent state
- ▶ comparison of S3T predictions with direct numerical simulations and verification of the theory
- ▶ stability of inhomogeneous turbulent states & relation with jet mergers
- ▶ relation of modulational instability of Rossby waves with S3T instability of homogeneous state (if time allows)
- ▶ summary

S3T generalizes the modulational instability of Rossby waves

MI is the hydrodynamic stability of finite amplitude Rossby waves (Lorenz 1972, Gill 1974, Connaughton et al. 2010)

$$\mathbf{p} = A \cos(\mathbf{p} \cdot \mathbf{x} - \omega_{\mathbf{p}} t)$$

We demonstrated that the problem of the hydrodynamic stability of **any** coherent nonlinear solution is mathematically equivalent to the S3T stability of the homogeneous equilibrium with the same eddy covariance spectrum.

$$\hat{C}^e(\mathbf{k}) = (2\pi)^2 p^4 |A|^2 [\delta(\mathbf{k} - \mathbf{p}) + \delta(\mathbf{k} + \mathbf{p})]$$

However, the two problems are very different:

MI studies the stability of infinitely coherent solutions

S3T studies the statistical stability of forced—dissipative flows with a given turbulence spectrum

Conclusions

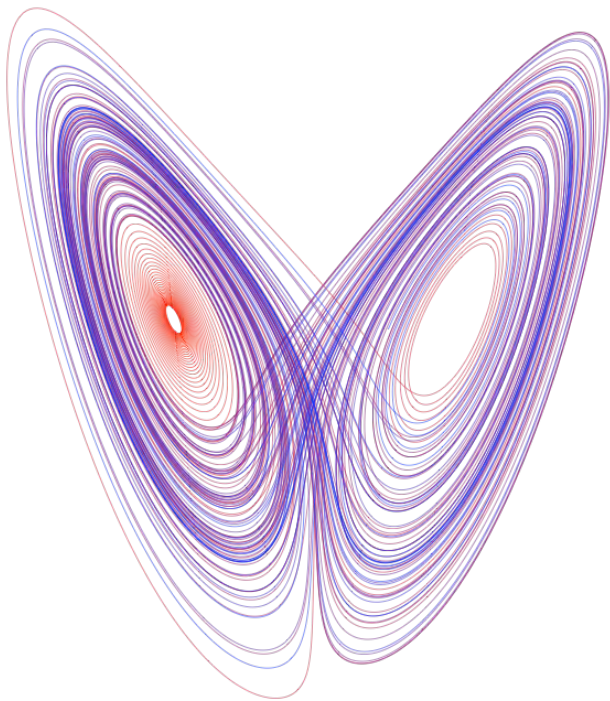
- ▶ S3T makes exact analytical predictions for the emergence of large-scale structure in planetary turbulence
- ▶ S3T predicts that the transition from a homogeneous to an inhomogeneous turbulent state occurs through a bifurcation of the statistical state dynamics
- ▶ S3T predicts the finite amplitude of the emergent large-scale structure
- ▶ The stability of inhomogeneous statistical turbulent equilibria (i.e. as in the Earth or Jupiter) can be studied within S3T framework and thus the sensitivity of the climate state of the planet can be determined

To understand turbulent flows one should adopt the perspective of the statistical state dynamics

Lorenz's vision



Ed Lorenz



“More than any other theoretical procedure, numerical integration is also subject to the criticism that it yields little insight into the problem. The computed numbers are not only processed like data but they look like data, and a study of them may be no more enlightening than a study of real meteorological observations. An alternative procedure which does not suffer this disadvantage consists of deriving *a new system of equations whose unknowns are the statistics themselves.*”

The Nature and Theory of the General Circulation of the Atmosphere,
by E. N. Lorenz, **1967**

S3T is a first step towards this *new system of equations*

References

- * Constantinou, Farrell & Ioannou (2014) Emergence and equilibration of jets in beta-plane turbulence: applications of Stochastic Structural Stability Theory. *J. Atmos. Sci.*, **71** (5), 1818–1842.
- * Bakas, Constantinou and Ioannou (2015) S3T stability of the homogeneous state of barotropic beta-plane turbulence, *J. Atmos. Sci.*, **71** (5), 1689–1712.
- * Constantinou (2015) Formation of large-scale structures by turbulence in rotating planets, Ph.D. thesis, University of Athens

ευχαριστώ
thanks